



A Bayesian approach for remote sensing of chlorophyll-*a* and associated retrieval uncertainty in oligotrophic and mesotrophic lakes

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ABSTRACT

Satellite remote sensing of chlorophyll-*a* concentration (chl_a) in oligotrophic and mesotrophic lakes faces uncertainties from sources such as atmospheric correction, complex inherent optical property compositions, and imperfect algorithmic retrieval. To improve chl_a estimation in oligo- and mesotrophic lakes, we developed Bayesian probabilistic neural networks (BNNs) for the Sentinel-3 Ocean and Land Cover Instrument (OLCI) and Sentinel-2 MultiSpectral Imager (MSI). The BNNs were built using an *in situ* dataset of oligo- and mesotrophic water bodies (1755 observations from 178 systems; median chl_a: 5.11 mg m⁻³, standard deviation: 10.76 mg m⁻³) and provide a per-pixel uncertainty percentage associated with retrieved chl_a. Shifts of oligo- and mesotrophic systems into the eutrophic regime, characterised by higher biomass levels, are widespread. To account for phytoplankton biomass fluctuation, a set of eutrophic lakes (167 observations from 31 systems) were included in this study (maximum chl_a 68 mg m⁻³). The BNNs were evaluated through five assessments including single day and time series match-ups with OLCI and MSI. OLCI BNN accuracy gains of >25% and MSI BNN accuracy gains of >15% were achieved in the assessments when compared to chl_a reference algorithms for oligotrophic waters (chl_a ≤ 8 mg m⁻³). In comparison to the reference algorithms, the accuracy gains of the BNNs decreased as chl_a and trophic levels increased. To measure the quality of the provided BNN uncertainty estimate, we calculated the prediction interval coverage probability (PICP), Sharpness and mean absolute calibration difference (MACD) metrics. The associated BNN chl_a uncertainty estimate included the reference *in situ* chl_a values for most observations (PICP ≥ 75%) across the different performance assessments. Further analysis showed that the BNN chl_a uncertainty estimate was not constantly well-calibrated across different evaluation strategies (Sharpness 1.7–6, MACD 0.04–0.25). BNN uncertainties were used to test two chl_a improvement strategies: 1) identifying and filtering uncertain chl_a estimates using scene-specific thresholds, and 2) selecting the most accurate prior atmospheric correction algorithm per individual satellite observation to retain chl_a with the lowest BNN uncertainty. Both strategies increased the quality of the chl_a result and demonstrated the significance of uncertainty estimation. This study serves as research on Bayesian machine learning for the estimation and visualisation of chl_a and associated retrieval uncertainty to develop harmonised products across OLCI and MSI for small and large oligo- and mesotrophic lakes.

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1. Introduction

The phytoplankton pigment chlorophyll-*a* concentration (chl_a) is widely used as a proxy of phytoplankton biomass and net primary production (Carlson, 1977; Huot et al., 2007; Poikane et al., 2010; Vörös and Padisák, 1991). Chl_a is thus an important indicator for the ecological integrity of aquatic ecosystems (Boyer et al., 2009). Optical sensors such as the MultiSpectral Imager (MSI) and Ocean and Land Cover Instrument (OLCI) aboard the Sentinel-2 and 3 satellites, respectively, are used to estimate chl_a and other optically active constituents (OACs) in lakes (Odermatt et al., 2018; Pahlevan et al., 2020; Toming et al., 2016). Oligotrophic and mesotrophic lakes have low to moderate levels of phytoplankton biomass (chl_a ≤ 25 mg m⁻³) and provide ecosystem services critical to human and animal welfare, such as drinking water, biodiversity, sediment retention and processing, fishery, and climate change mitigation (Grizzetti et al., 2016; Sterner et al., 2020).

Spectral remote sensing reflectance ($R_{rs}(\lambda)$) contains information about water column properties and components such as phytoplankton (Mobley, 1999; O'Reilly et al., 1998). Many chl_a estimation algorithms relate optical features in the red and near-infrared (NIR) areas of the reflectance spectrum to chl_a (Odermatt et al., 2012). For example, the contribution from phytoplankton backscattering to the reflectance spectrum forms distinct reflectance peaks near 560 and 700 nm (Suits, 1975). Further, a phytoplankton pigment feature at 675 nm can be offset against the NIR scattering peak around 700 nm (Gitelson, 1992; Mittenzwey et al., 1992). A variety of two, three and four-band combination algorithms were developed over the last decades based on the 700 and 665 nm band ratio to capture the red absorption signature of chl_a (Dall'Olmo et al., 2003; Gilerson et al., 2010; Gons, 1999; Gurlin et al., 2011; Mishra and Mishra, 2012; Moses et al., 2009; Pepe et al., 2001). Chl_a absorption and fluorescence peaks in the 665–685 nm region are also at the basis of line height algorithms such as the fluorescence line height (FLH) (Gower, 1980; Gower et al., 2004; Gupana et al., 2021), maximum peak height (MPH) (Matthews et al., 2012; Matthews and Odermatt, 2015) and maximum chlorophyll index (MCI) (Binding et al., 2013; Gower et al., 2005).

In recent decades, the majority of inland water quality studies have concentrated on eutrophic water bodies (Clark et al., 2017; Coffey et al., 2021; Dekker and Peters, 1993; Simis et al., 2005; Urquhart et al., 2017). In comparison, oligo- and mesotrophic water bodies are under-sampled globally (Filazzola et al., 2020), limiting the development and validation of designated chl_a estimation methods. Furthermore, recent research indicates that chl_a estimation in oligo- and mesotrophic systems is associated with higher uncertainties than in eutrophic conditions (Liu et al., 2021; Mouw et al., 2013; Neil et al., 2019; Werther et al., 2022). The causes of estimation uncertainty are manifold. Absorption by coloured dissolved organic matter ($a_{CDOM}(\lambda)$) and non-pigmented particulate absorption ($a_{NAP}(\lambda)$) may be larger than by phytoplankton pigments ($a_p(\lambda)$), particularly at short wavelengths where band arithmetic algorithm specificity then breaks down (Härmä et al., 2001; Kutser et al., 2016; Mouw et al., 2013; Neil et al., 2019; Schalles, 2006). For analytical algorithms that invert a reflectance spectrum into inherent optical properties (IOPs), $a_p(\lambda)$ must be precisely estimated to relate it to chl_a (Werdell et al., 2018). Effects such as pigment packaging impact the linearity of the relationship between chl_a and $a_p(\lambda)$ (Grzymalski et al., 1997; Kirk, 1994; McKee et al., 2014). The scaling of $a_p(\lambda)$ to chl_a further depends on phytoplankton type and particle size (Johnsen et al., 1994; Lutz, 2001; Simis et al., 2007). Therefore, for analytical algorithm application to estimate chl_a, prior information about $a_p(\lambda)$ variability is required. Without system-specific calibration, the underlying bio-optical models use default pigment mass-specific phytoplankton absorption ($a_p^*(\lambda)$) values. $a_p^*(\lambda)$ natural variability spans an order of 4 magnitudes in the ocean (Bricaud et al., 1995, 1998, 2004), and its variability is unknown for global lakes, yet subject to measurement uncertainty (McKee et al., 2014). For algorithms that focus on the red/NIR spectral area, the phytoplankton absorption peak at 675 nm in lakes may be undetectable in conditions where CDOM and NAP absorption is high (Kutser et al., 2016). For oligotrophic

lakes, the red-NIR area of the spectrum may have low signal to noise ratios, causing chl_a-related peaks to be below the detection limit.

In recent years, machine learning (ML) methods have been developed to estimate chl_a in inland water bodies. Neural networks (NNs) developed for lakes overcome some of the issues encountered with band ratio methods and analytical approaches when assessed over large chl_a ranges (Pahlevan et al., 2020, 2022; Schaeffer et al., 2022; Smith et al., 2021). For a ML approach to accurately handle unseen observations, large training datasets of *in situ* and/or simulated measurements are necessary. Because the spatial and temporal scope of the training set is frequently constrained, ML techniques are prone to produce estimation errors when a so-called dataset shift happens. A dataset shift occurs when the IOPs and OACs of a particular, unknown input deviate considerably from the measurements utilised during ML model training (Moreno-Torres et al., 2012; Ovadia et al., 2019). In lake remote sensing, chl_a estimation uncertainty generated by a dataset shift or other factors such as preceding AC and surrounding land effects on the radiance distribution are typical. Current algorithms designed for lakes, however, do not account for the uncertainty associated with estimated chl_a. Because chl_a estimation uncertainty is larger in oligo- and mesotrophic lakes than in eutrophic lakes (Neil et al., 2019), a provision of confidence is a desirable feature to enable the detection and handling of questionable chl_a estimates.

Bayesian probability theory offers a mathematical tool to reason about uncertainty (Ghahramani, 2015). Bayesian probabilistic reasoning applied to NNs usually comes at a prohibitive computational cost. A NN with distributions placed over the weights has long been studied as a Bayesian Neural Network (BNN) (MacKay, 1992; Neal, 1996) commonly through variational inference (VI) (Jordan et al., 1999; Saul et al., 1996), but with limited success (Gal and Ghahramani, 2016; Graves, 2011). An approach based on VI brings higher model complexity, which for the field of aquatic remote sensing is prohibitive: to represent uncertainty, the number of parameters in BNNs based on VI is doubled when compared to the same non-Bayesian NN size. To train a network with millions of parameters large input datasets are required. However, for inland aquatic remote sensing large datasets are sparse and only just mature through community-wide efforts to collate datasets ready for use with recent ML approaches (Cao et al., 2020; Pahlevan et al., 2022; Werther et al., 2021).

Here, we investigate a potential solution in BNNs developed through Monte Carlo dropout (Gal and Ghahramani, 2016) to obtain chl_a and a well-calibrated uncertainty estimate from a limited set of input training samples. Monte Carlo (MC) dropout refers to the process of randomly turning off all outgoing connections from a neuron in a NN through a previously determined probability p (Srivastava et al., 2014). Dropout can be activated during the application to unknown observations, which results in an ensemble of NNs to estimate a final chl_a value: for a single input observation, MC dropout generates a previously defined number (S) of NN variants, each with a different set of activated neurons during prediction, producing S unique chl_a estimates. The S estimates form a predictive distribution which is then averaged to obtain a final chl_a value.

Based on the largest available *in situ* dataset covering typical oligo- and mesotrophic lake optical properties, we aim to: 1) improve the chl_a estimation accuracy via OLCI and MSI sensors over both small and large oligo- and mesotrophic lakes through Bayesian probabilistic machine learning, 2) provide a well-calibrated, per-pixel uncertainty percentage associated with estimated chl_a, and 3) demonstrate how the provided uncertainty can be used to deal with dataset shifts and AC errors commonly affecting satellite remote sensing of chl_a.

2. Datasets

2.1. Development dataset

A dataset of 1755 *in situ* observations from 178 lakes and reservoirs was compiled from six sources to develop and evaluate the BNNs: (i) LIMNADES (Lake Bio-optical Measurements and Matchup Data for Remote Sensing (<https://limnades.stir.ac.uk/>)), (ii) Wisconsin DNR (U.S.), (iii)

Earth and Planetary Observation Sciences (EPOS) (Scotland, UK), (iv) the University of Tartu (Estonia), (v) the University of Waikato (New Zealand (NZ)) and the (vi) LÉXPLORE Platform (Lake Geneva, Switzerland). Only samples containing above-water $R_{rs}(\lambda)$ and chl_a measurements were considered for this study. Where available, total suspended matter (TSM) dry weight and the absorption by CDOM at 443 nm ($a_{CDOM}(443)$) were included in the dataset. For algorithm development and validation we separated the measurements from these datasets into 10 regions (Fig. 1) to evaluate the BNNs on independent geographical regions, as previously demonstrated for ML algorithm development (O'Shea et al., 2021; Pahlevan et al., 2022). Details about the regions and included water bodies are provided in Appendix 1. Datasets in LIMNADES were described in Spyraikos et al. (2018) and their original publications (Binding et al., 2008, 2010; Bradt, 2012; Bresciani et al., 2011; Dall'Olmo et al., 2005; Giardino et al., 2005, 2013, 2014, 2015; Gitelson et al., 2007, 2008; Gons et al., 2008; Guanter et al., 2010; Li et al., 2013, 2015; Manzo et al., 2015; Matsushita et al., 2015; Matthews and Bernard, 2013; Moore et al., 2014; Ruiz-Verdú et al., 2005, 2008; Schalles and Hladik, 2012; Wang et al., 2018; Yacobi et al., 2011). Measurements from Wisconsin DNR were taken across the U.S. State of Wisconsin between spring 2014 and autumn 2016. The Wisconsin DNR dataset was used in recent algorithm development and inter-comparison studies (Pahlevan et al., 2020, 2021b). Collection methods and details are provided in Werther et al. (2022). Measurements provided by the University of Tartu were previously detailed in Kutser et al. (2013); Soomets et al. (2020). The measurements from lakes in NZ refer collectively to two datasets. The first one was collected by the University of Waikato between 2015 and 2019 and has been previously published (Pahlevan et al., 2020, 2022). The second dataset refers to data collected jointly by the University of Waikato and University of Stirling in 2020 (referred to as NZ 2020). The NZ 2020 dataset is unpublished and described in Appendix 2. Surface chl_a obtained through calibration of *in vivo* fluorescence measurements and $R_{rs}(\lambda)$ from the LÉXPLORE platform (<https://lexplore.info/>) in Lake Geneva were measured between 2018/10 and 2021/09. Measurement details are provided in Appendix 3.

Optical Water Types (OWTs) of Spyraikos et al. (2018) were used to define the development and application range of the BNNs (Fig. 2). OWT

membership was derived through estimation of the spectral angle (Kruse et al., 1993; Liu et al., 2021). For all $R_{rs}(\lambda)$ an OWT membership score was calculated and the OWT with the highest similarity score assigned. Observations assigned to OWT 5 were small ($n = 40$) in this dataset. The corresponding systems were high in CDOM absorption, TSM and low in chl_a, known to be challenging for retrieval algorithms (Kallio et al., 2015; Kutser et al., 2016; Toming et al., 2016; Werther et al., 2022).

Phytoplankton productivity in lakes is rarely steady and changes with season and external forcings such as wind (Anneville et al., 2004; Rusak et al., 2018). Lakes that are oligotrophic might suddenly become eutrophic and vice versa (Gilarranz et al., 2022; Minaudo et al., 2021). To account for changes in phytoplankton biomass, the collected data was focused on oligo- and mesotrophic lakes, with an expansion to eutrophic systems (chl_a median 5.11 mg m⁻³, mean 9.45 mg m⁻³, standard deviation 10.76 mg m⁻³) (Fig. 3). 167 observations (9.5%) from 31 systems in the dataset exhibit chl_a > 25 mg m⁻³ with a maximum of 68 mg m⁻³. In this study, we adopt the OECD (1982) trophic status scheme, which serves as an international standard for assessing the ecological status of lakes (Carvalho et al., 2013). OECD (1982) classified water trophic status as oligotrophic for chl_a ≤ 8 mg m⁻³, mesotrophic until 25 mg m⁻³, and eutrophic thereafter. We also refer to oligo- and mesotrophic systems as low - moderate biomass lakes until 25 mg m⁻³, because eutrophic lakes can quickly reach phytoplankton chl_a > 100 mg m⁻³ (Matthews, 2014; Matthews et al., 2012; Neil et al., 2019; Pahlevan et al., 2020).

2.2. Satellite data processing

Match-ups were generated for the period 2018/10 to 2021/09 between coinciding *in situ* measurements taken on the LÉXPLORE platform (Lake Geneva) and satellite observations from Sentinel-3A/B (S3 A/B) OLCI 300 m resolution and Sentinel-2A/B (S2 A/B) MSI 20 m resolution. Surface *in situ* chl_a from a Thetis profiler was obtained between 9 and 12.30 a.m. (UTC) (see Appendix 3 for details). OLCI 3A and B overpassed the platform between 10 and 11.30 a.m. (UTC). To generate the match-up, the nearest OLCI and MSI sensor overpass within ± 3 hours of the *in situ* Thetis chl_a measurement was chosen. The S2 A/B overpass

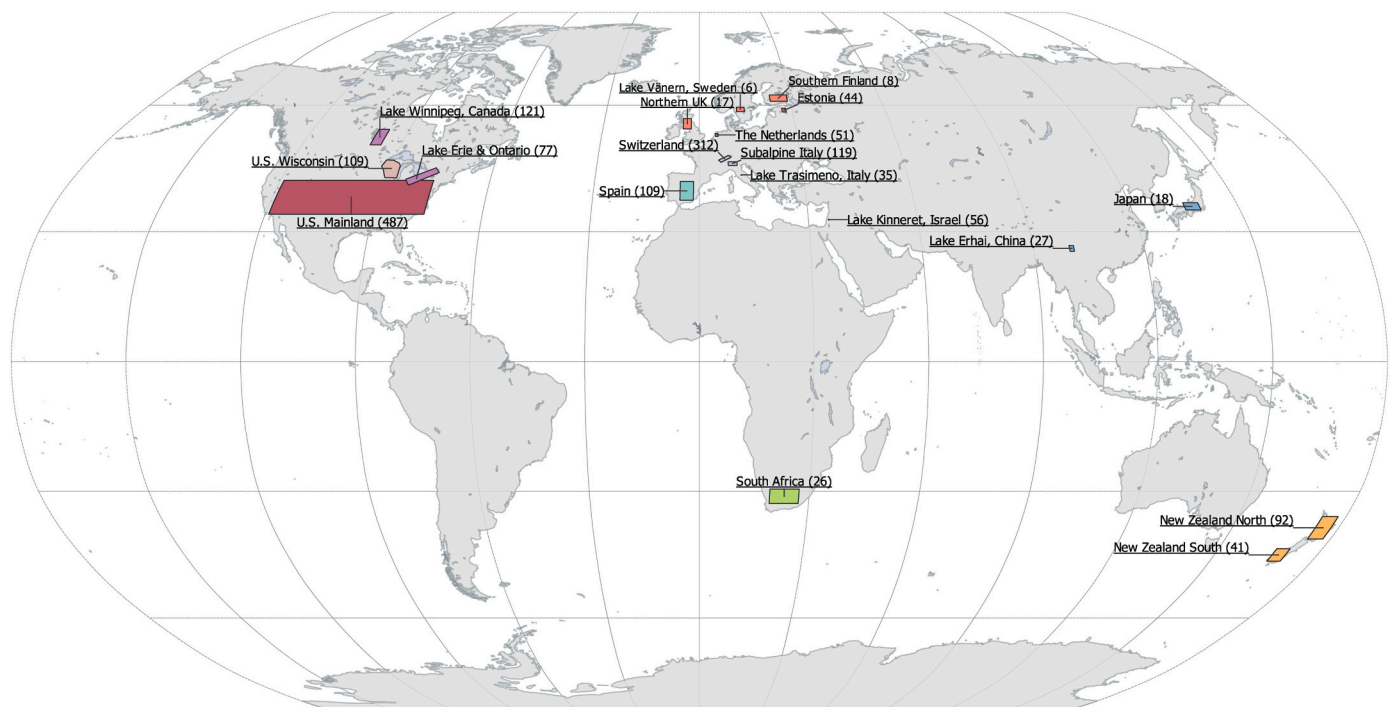


Fig. 1. Spatial distribution of the 10 *in situ* measurement regions constituting the dataset of this study. Grouped regions share the same colour. Number of samples taken in the indicated polygon is shown in brackets. See Appendix 1 for a detailed description of the dataset regions.

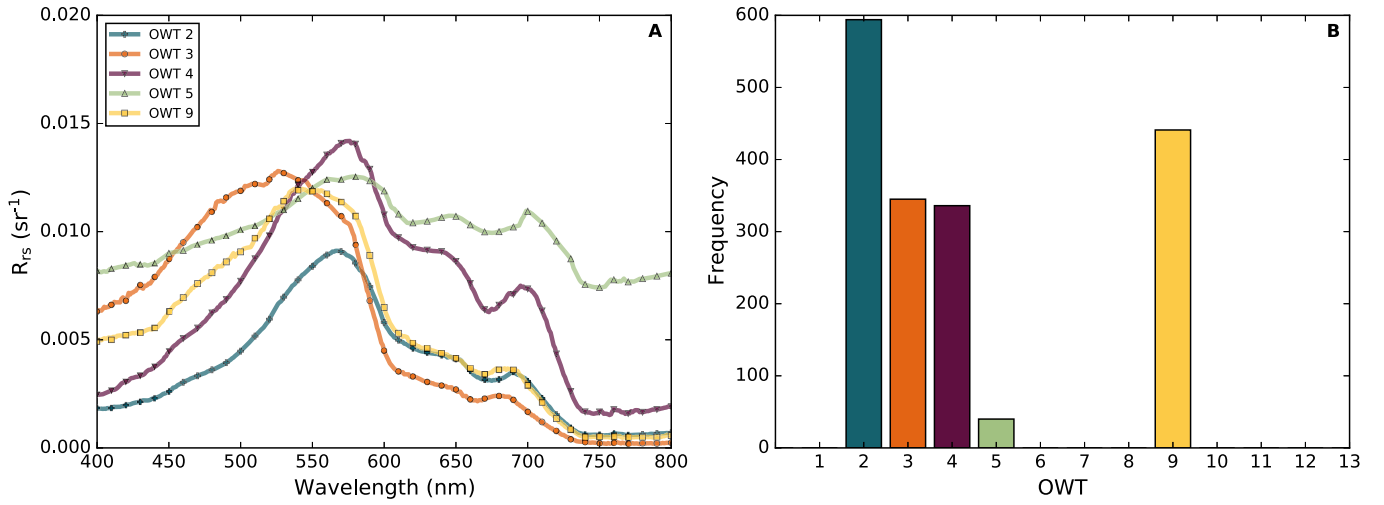


Fig. 2. OWTs 2, 3, 4, 5 and 9 used for the development and application of the BNNs. (A) Spectral medians. (B) Frequency per OWT.

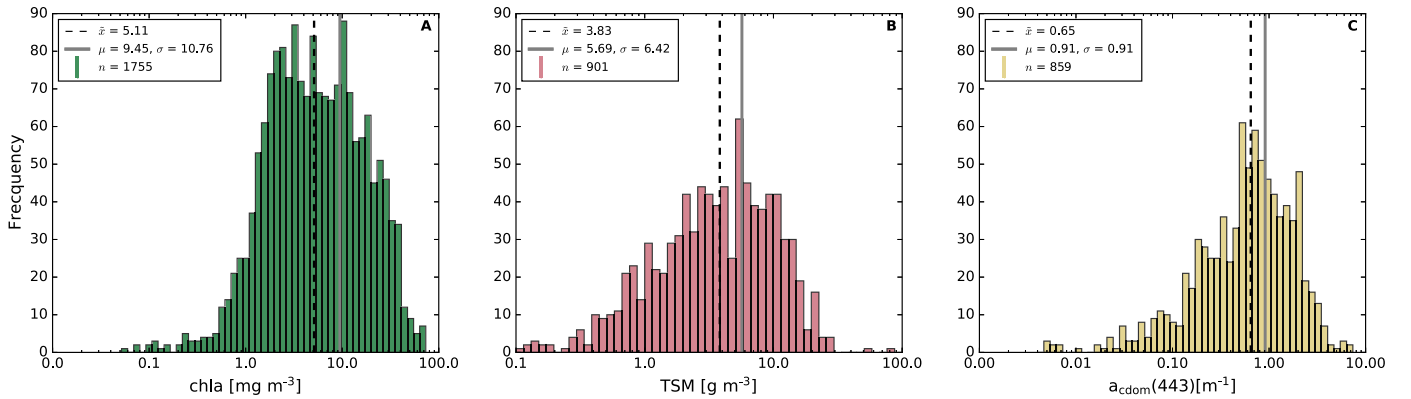


Fig. 3. Log-distributions of (A) chl a, (B) TSM and (C) $a_{CDOM}(443)$ for the entire dataset. Denoted are median (\hat{x}), mean (μ) and standard deviation (σ) of the respective parameter.

frequency over LÉXPLORE is one-fifth of the S3 A/B OLCI frequency but occurs in the same time frame. For OLCI and MSI (in their respective spatial resolution) valid pixels south of the platform were identified through the IdePix (Identification of Pixel features) algorithm (Skakun et al., 2022; Wevers et al., 2021). The centre pixel of 3×3 valid pixels was used as the match-up location. Data flagged by IdePix as invalid, cloud (including ambiguous, sure, and a 2-pixel buffer) or cloud shadow, snow/ice, bright, coastline land, white and glint risk were excluded.

For AC of OLCI and MSI data, the POLYMER (Steinmetz et al., 2011) and Case 2 Region CoastColour (C2RCC) (Brockmann et al., 2016) algorithms were selected. POLYMER and C2RCC AC performances were previously compared for OLCI and MSI and found to perform well for the OWTs included in this study (Liu et al., 2021; Pahlevan et al., 2021a; Warren et al., 2019, 2021). Both ACs are widely used in combination with algorithms for the retrieval of chl a in lakes (Kravitz et al., 2020; Liu et al., 2021; O'Shea et al., 2021; Pahlevan et al., 2020, 2022; Pereira-Sandoval et al., 2019; Smith et al., 2021). A comparison of $R_{rs}(\lambda)$ obtained from the LÉXPLORE platform versus OLCI/MSI POLYMER/C2RCC is provided in Appendix 4.

3. Methodology

3.1. Bayesian Neural Networks

At the heart of all BNNs is the Bayesian theorem:

$$p(\theta|D) = \frac{p(D|\theta)p(\theta)}{p(D)}, \quad (1)$$

where $p(\theta|D)$ is the posterior, i.e., the probability of a certain value of a neural network parameter θ given the data D , $p(D|\theta)$ the likelihood, $p(\theta)$ the prior and $p(D)$ the quantity (also known as the marginal likelihood) (Bolstad and Curran, 2016). In Bayesian statistics, θ is not defined by one value but has an uncertainty described by a probability distribution $p(\theta)$. This distribution, $p(\theta)$, defines the probability value of each parameter value θ . For a Bayesian model to obtain a predictive chl a distribution, the weight distribution of the neural network is averaged:

$$p(y|x_{test}, D) = \sum_i p(y|x_{test}, \theta_i) \bullet p(\theta_i|D), \quad (2)$$

where x_{test} is an input test $R_{rs}(\lambda)$ observation and θ_i denote the weights w_i of the NN. Replacement of NN weights with distributions is computationally costly and requires large amounts of training data, which is scarce in lake remote sensing. In this study, we therefore used Monte Carlo (MC) dropout applied to NN layers (Gal and Ghahramani, 2016). MC dropout replaces each fixed θ_i of a deterministic NN with a binary distribution, whereby either zero or the value θ_c for a NN connection are obtained. Using a dropout chance of 25%, the NNs estimated S -times chl a for a single input reflectance spectrum. Each of the S estimates originated from a different NN variant that corresponded to a sampled network constellation. Combining the S dropout estimates resulted in a chl a probability distribution:

$$p(y|x_{test}, D) = \frac{1}{S} \sum_{i=1}^S p(y|x_{test}, \theta_i), \quad (3)$$

which is an empirical approximation to Eq. 2 that captured both the NN model (epistemic) and the data intrinsic aleatoric uncertainty (Abdar et al., 2021). S is a user-supplied value. The higher S , the more NN dropout variants and chla estimates are generated to form the probability distribution. We chose 50 in this study as a trade-off between required computation time and performance. We then sampled from all determined Gaussian distributions: $y \sim N(\mu_{x,\theta}, \sigma_{x,\theta})$, where μ_{θ} is the chla mean and σ_{θ} the standard deviation (Fig. 4). Because σ was a distribution, for each chla estimate y we calculated the 95% confidence interval (CI) and then estimated the width of the CI (CI_w) to get an absolute number. CI_w was divided by two to obtain uncertainty \pm of y and the entire term was multiplied by 100 to obtain a BNN percentage uncertainty for estimated chla:

$$\varepsilon(\%) = \left(\frac{CI_w}{2} \right) \cdot \frac{100}{y}. \quad (4)$$

Determination of the BNN hyper-parameters, such as the probability that a neuron was dropped, is explained in section 3.3.1.

3.2. BNN processing

In situ $R_{rs}(\lambda)$ were convolved to the relative spectral response (RSR) of the OLCI instruments on Sentinel-3A and 3B and the MSI instruments on Sentinel-2A and B, and were used as inputs to the BNNs. Distinct BNNs were developed for OLCI S3A/B and MSI S2A/B. For the OLCI BNNs we selected the 12 bands from 413 through 778 nm (omitting the oxygen bands at 761, 764 and 767 nm). The waveband centred at 400 nm was excluded because the majority of *in situ* $R_{rs}(\lambda)$, did not extend to <400 nm and therefore could not be convolved to match the OLCI band. For MSI, all 7 bands from 443 to 783 nm were selected.

For oligotrophic waters with chla <5 mg m⁻³, negative $R_{rs}(\lambda)$ in one or more red-NIR wavebands may occur because of low signal to noise ratios. Negative values in these wavelengths can cause unpredictable behavior of

the BNNs that may lead to high chla estimation uncertainty. To enable the BNNs to produce chla estimates even when negative values in red-NIR wavelengths occur, all negative $R_{rs}(\lambda)$ values ≥ 665 nm in the *in situ* dataset were set to zero for BNN training and during pre-processing of test observations (including AC-derived $R_{rs}(\lambda)$). $R_{rs}(\lambda)$ features were then normalized to fall between the 0–1 range. Each feature was treated individually such that it was in the given range of the dataset between 0 and 1. Normalisation reduced the required training time and smoothed the process of minimising the loss function of the network, the negative log-likelihood (NLL) in this study. NLL is a standard loss function to measure the quality of a probabilistic model outcome (Hastie et al., 2001).

Since the BNNs are based on MC dropout, a stochastic process is involved. To obtain a representative BNN for the different evaluation strategies (see section 3.3 below), 10 BNNs were trained for each sensor and applied to the respective test observations once. The median of the 10 different networks was used as the final estimate we report, following recent practise of NNs with stochastic elements (O'Shea et al., 2021; Pahlevan et al., 2020, 2022). For the evaluation on *in situ* data, the S3A and S2A BNNs versions were used. For match-ups with MSI 2A/B and OLCI 3A/B the BNN version corresponding to the respective sensor was selected.

3.3. BNN performance assessment

Five different assessments were conducted to evaluate the quality of the chla and uncertainty BNN products under varying conditions: (i) a randomised 50/50 training/test data split, (ii) a leave-one-out exercise (LOO) using *in situ* observations, (iii) BNN OLCI and MSI LÉXPLORE match-ups, (iv) a LÉXPLORE 3 ½ year time-series (including the match-ups) and lastly (v) single-day product visualisations over both large and small lakes in Europe, New Zealand, Africa and Canada. The BNNs were compared to state-of-the-art chla algorithms. Details of these assessments are provided in the following sections.

3.3.1. 50/50 training/test data split

For an assessment of overall BNN performance on *in situ* data, the development dataset ($n = 1755$) was split into 50% training ($n = 878$) and 50% test ($n = 877$) sets, following recent ML assessment practises (O'Shea et al., 2021; Pahlevan et al., 2020). The observations in the two sets were

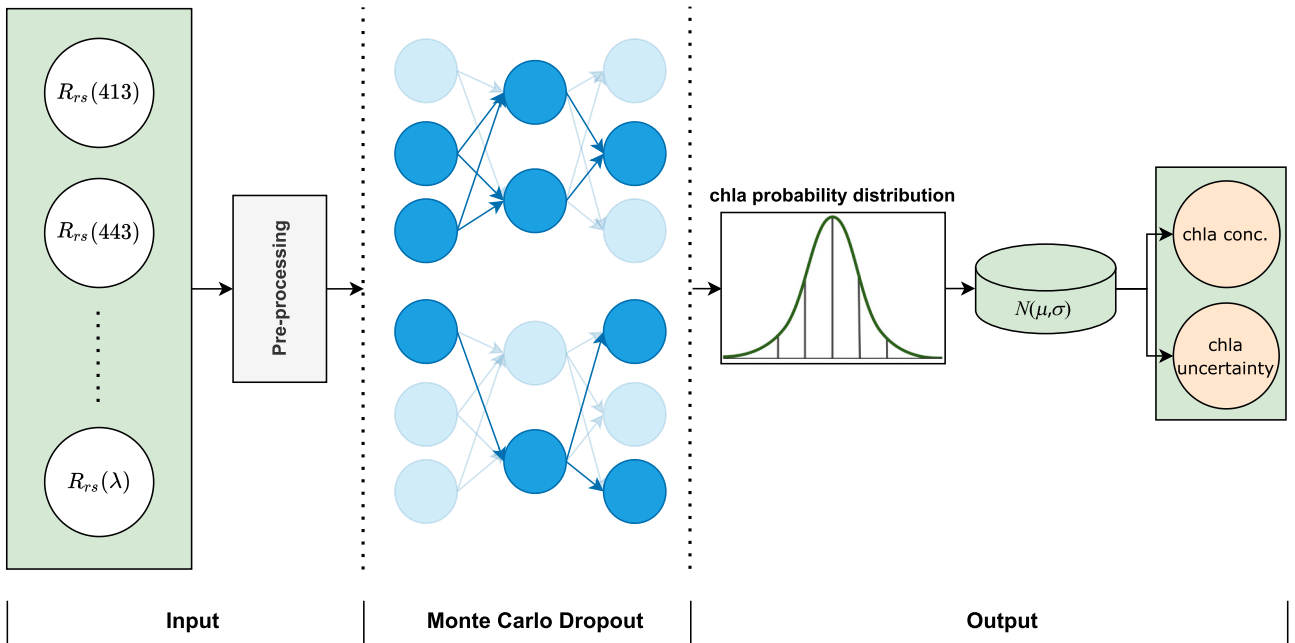


Fig. 4. Processing scheme of the BNNs based on Monte Carlo dropout. For MSI the first band is 443 nm. Pre-processing includes the spectral convolution of the training data, normalisation of both training and unknown (test) observations and treatment of negative $R_{rs}(\lambda)$ values.

randomly drawn from all regions of the entire dataset. The 50% training data were further split into a training (60%; 526 observations) and validation set (40%; 352 observations). An initial OLCI and MSI BNN was constructed using this training subset ($n = 526$), and the performance of these BNN versions was evaluated on the validation set ($n = 352$) to find an optimal architecture and to tune the hyper-parameters of the BNNs through Bayesian optimisation implemented in the Python package *Weights & Biases* (Biewald, 2020; Werther et al., 2021). BNNs were continually improved by tracking the loss on the validation data. Once the validation loss did not decrease further, the evaluation on the validation set was concluded. The BNNs were then re-trained with the entire training set (878 / 1755 observations) and subsequently applied to the test set using the previously determined architecture and hyper-parameters. The hyper-parameters identified in this 50/50 assessment were used across the other performance assessments to allow comparisons of the same model architecture with different training and test sets. Furthermore, the OLCI and MSI BNNs distributed in conjunction with this article conform to the single hyper-parameter configuration utilised throughout the assessments. The OLCI and MSI BNN performances on the test set across OWTs and TS classes were tabulated.

A randomised 50/50 split of training and test data has two major limitations. First, training and test datasets share the same distribution of OACs (assured through prior randomisation of the overall dataset), which may not occur in a situation outside of model development. Performance under dataset shifts can therefore not be analysed. Second, the performance of the BNN OLCI and MSI versions in this assessment does not represent their final models, as these versions were only trained with 50% of the training data. The final BNN versions made publicly available with this article were re-trained with the entire dataset ($n = 1755$). The final hyper-parameter configuration of the OLCI and MSI BNNs is listed in Appendix 5.

3.3.2. Leave-one-out

We simulated dataset shifts with all available *in situ* measurements through a leave-one-out (LOO) strategy (O'Shea et al., 2021; Pahlevan et al., 2022). In LOO, a BNN was trained with measurements from all regions except for one region. The BNN was then applied to the left-out region and the entire process was repeated until each region was left-out. Through the LOO strategy we assess the generalisation performance of the BNN for independent systems (represented by a region) that were not part of the training process, and which may thus not share the same OAC distribution. We note that the performance of the BNNs for different regions may underestimate overall model performance. Whilst different *in situ* measurement techniques between the datasets were used, the measurement consistency between regions is unknown. Individual regions may thus carry varying degrees of measurement uncertainty.

3.3.3. Lake Geneva match-ups and time series assessments

To evaluate the BNNs with independent satellite match-ups, a training set excluding observations from Lake Geneva was created. OLCI and MSI BNNs were trained with the exact same set of bands and architecture as in the 50/50 training/test split and LOO assessments. The match-up evaluation on Lake Geneva facilitated to assess how well the BNNs trained with a semi-global *in situ* dataset of lake properties transfer to an individual system measured through satellite sensors, thereby representing a common use-case of the newly developed algorithm.

3.3.4. Comparison with other chl_a algorithms

Using the 50/50 split, LOO and match-up performance assessments we evaluated the accuracy of the OLCI and MSI BNNs *versus* five state-of-the-art algorithms designed for large chl_a concentration ranges (hereafter reference algorithms). The five algorithms are the OLCI and MSI Mixture Density Networks (MDNs) (Pahlevan et al., 2020), the red/NIR semi-analytical Gons05 (Gons et al., 2005), the red/NIR band ratio G11 (Gurlin et al., 2011), the blue/green band ratio OC3 (O'Reilly et al., 1998; O'Reilly and Werdell, 2019), and the Blend algorithm (Smith

et al., 2018) which switches between OCI (Hu et al., 2012) and the red/NIR semi-analytical approach by Gilerson et al. (2010), that blends their estimates using empirically derived thresholds.

To reduce bias resulting from calibration of the reference algorithms on different datasets, the G11, OC3 and Blend model coefficients were optimised (denoted as -opt) through non-linear least squares fitting against the same training datasets used in the evaluation strategies (identical to the observations for training the BNNs). We used the original OLCI and MSI MDN versions that were designed for general applicability to inland waters (including low - moderate biomass lake conditions). For the OLCI and MSI MDNs specifically, in the LOO assessment regions 5 and 9 were excluded from the evaluation because these were already used in full as part of the training set of said algorithms. Their inclusion would thus not represent an independent comparison per region. Gons05 optimisation was not undertaken as it requires measurements of phytoplankton mass-specific absorption at 665 nm ($a_{\phi}^*(665)$) and backscatter, which the collated dataset did not sufficiently include.

Algorithms using OLCI wavebands were included in their originally published configuration (denoted as -org) and with optimised coefficients (denoted as -opt). For Blend in the OLCI configuration only the OC4 algorithm was optimised, and for Blend in the MSI configuration the OC3 algorithm was used instead of OC4 because the 510 nm band in OC4 is not available on MSI. MSI band positions of Gons05, G11 and OC3 (also used in Blend) were slightly shifted, namely from 490, 708 and 778 nm to 492, 704 and 783 nm, which was successfully demonstrated previously in Warren et al. (2021). Optimised model coefficients are tabulated in Appendix 6. For the Lake Geneva match-ups, all measurements were provided for optimisation, except for the observations measured in Lake Geneva.

3.4. Accuracy and uncertainty calibration metrics

Common performance metrics were calculated to assess the chl_a retrieval accuracy between algorithm estimated (e) and observed (o) *in situ* chl_a: the median symmetric accuracy (MdSA, in %), the symmetric signed percentage bias (Bias, in %), the mean absolute difference (MAD) and median absolute percentage difference (MAPD, in %) (Morley et al., 2018; Seegers et al., 2018, 2021; Smith et al., 2021):

$$MdSA : \left(10^{\text{median} \left(\left| \log_{10} \left(\frac{e}{o} \right) \right| \right)} - 1 \right) \times 100. \quad (5)$$

$$Bias : \left(10^{\text{median} \left(\log_{10} \left(\frac{e}{o} \right) \right)} - 1 \right) \times \text{sign}(MR) \times 100, \quad (6)$$

where $\text{median ratio } (MR) = \text{median} \left(\log_{10} \left(\frac{e}{o} \right) \right)$.

$$MAD : 10^Z - 1, Z = \left[\frac{\sum_{i=1}^n \left| \log_{10}(e_i) - \log_{10}(o_i) \right|}{N} \right]. \quad (7)$$

$$MAPD : \text{median} \left(\left| \frac{e_i - o_i}{o_i} \right| \right) \times 100 \text{ for } i = 1, \dots, N. \quad (8)$$

BNN uncertainty estimates may not capture the true data distribution (Lakshminarayanan et al., 2017). For example, a 90% confidence interval (CI) may not contain the reference *in situ* chl_a concentration in 9/10 scenarios. The BNN uncertainty estimate is miscalibrated when the CI does not include the reference *in situ* chl_a. To assess the quality of the BNN chl_a uncertainty calibration we calculated three metrics between the BNN estimated (e) and observed (o) *in situ* chl_a:

1. Prediction interval coverage probability (PICP; denoted ρ). The percentage of observations for which *in situ* chl_a lies within the 95% confidence interval of the BNN chl_a estimate:

$$\rho = \frac{1}{N} \sum_{n=1}^N \mathbb{1}_{o_n \leq e_n^{high}} \bullet \mathbb{1}_{o_n \geq e_n^{low}}, \quad (9)$$

where e_n^{high} is the 97.5% percentile and e_n^{low} is the 2.5% percentile of the BNN estimated chl_a for an input x_n (Yao et al., 2019). The higher the ρ [%] value, the more *in situ* chl_a values were covered by the BNN uncertainty estimate.

2. Sharpness (denoted σ). The standard deviation (var) of a chl_a estimate whose cumulative distribution function (F_N) should be small:

$$\sigma = \sqrt{\frac{1}{N} \sum_{n=1}^N \text{var}(F_n)^2}. \quad (10)$$

Sharpness represents the average of the BNN estimated chl_a standard deviations (Tran et al., 2020).

3. Mean absolute calibration difference (MACD; denoted τ). MACD is a statistic to measure calibration relative to an ideal reliability diagram. Reliability diagrams show expected observation accuracy as a function of confidence (Degroot and Fienberg, 1983; Niculescu-Mizil and Caruana, 2005). To compute MACD, chl_a estimates were sorted and divided into B equally-spaced bins ($B = 100$ in this study) with an approximately equal number of BNN chl_a estimates in each bin:

$$\tau = \sum_{b=1}^B \frac{n_b}{N} \bullet |acc(b) - conf(b)|, \quad (11)$$

where n_b is the number of chl_a estimates in bin b , N is the total number of observations, and $acc(b)$ and $conf(b)$ are the accuracy and confidence in bin b (Guo et al., 2017; Nixon et al., 2019). The smaller the values of Sharpness and MACD, the more accurate the uncertainty calibration.

4. Results

4.1. 50/50 dataset split

OLCI and MSI BNNs outperformed the reference chl_a algorithms (MdSA difference > 18%) (Figs. 5, 6). The OLCI BNN was slightly more accurate than its MSI version (1.44% MdSA difference). In the employment of the reference chl_a algorithms, lower accuracy in the estimation of chl_a from MSI compared to OLCI was also observed (Fig. 6). For OLCI, optimisation of model coefficients from G11, OC3 and Blend improved the estimates by 35–75% (MdSA) for oligotrophic waters and 15–35% (MdSA) for mesotrophic waters (Table 1). For eutrophic lakes the original algorithm configurations were more accurate, because the dataset of this study used to optimise the algorithms only included a small set of high biomass observations. The coefficients of the original formulations were obtained from larger eutrophic datasets and are thus more applicable. The optimisation process of the coefficients is considered transferable to MSI. Large over- and underestimates of chl_a were observed for concentrations between 0 and 1 mg m⁻³ across all the algorithms. Only 65 data points (7.4% of 877 test observations) were available in this concentration range. Chl_a estimates from the algorithms used for comparison became markedly more accurate for concentrations >10 mg m⁻³. Varying retrieval accuracies became more apparent when the assessment was tabulated per TS class (Table 1) as the OLCI BNN was >45% and the MSI BNN > 38% more accurate for oligotrophic waters (chl_a 0–8 mg m⁻³) than the reference algorithms. Red/NIR algorithms started to become highly accurate in mesotrophic waters (chl_a 8–25 mg m⁻³) (Table 1). For eutrophic waters (chl_a > 25 mg m⁻³) the red/NIR OLCI Gons05 algorithm was most accurate (MdSA 13.98%), but its performance did not transfer to MSI. Across OWTs, the BNNs consistently outperformed the reference algorithms (Table 2). For OWT 5 it is to note that only 13 observations were assigned to it, thus the evaluation is not representative. The low number in this OWT originated from the randomisation of the entire dataset prior to training and evaluation of the algorithms. Uncertainty

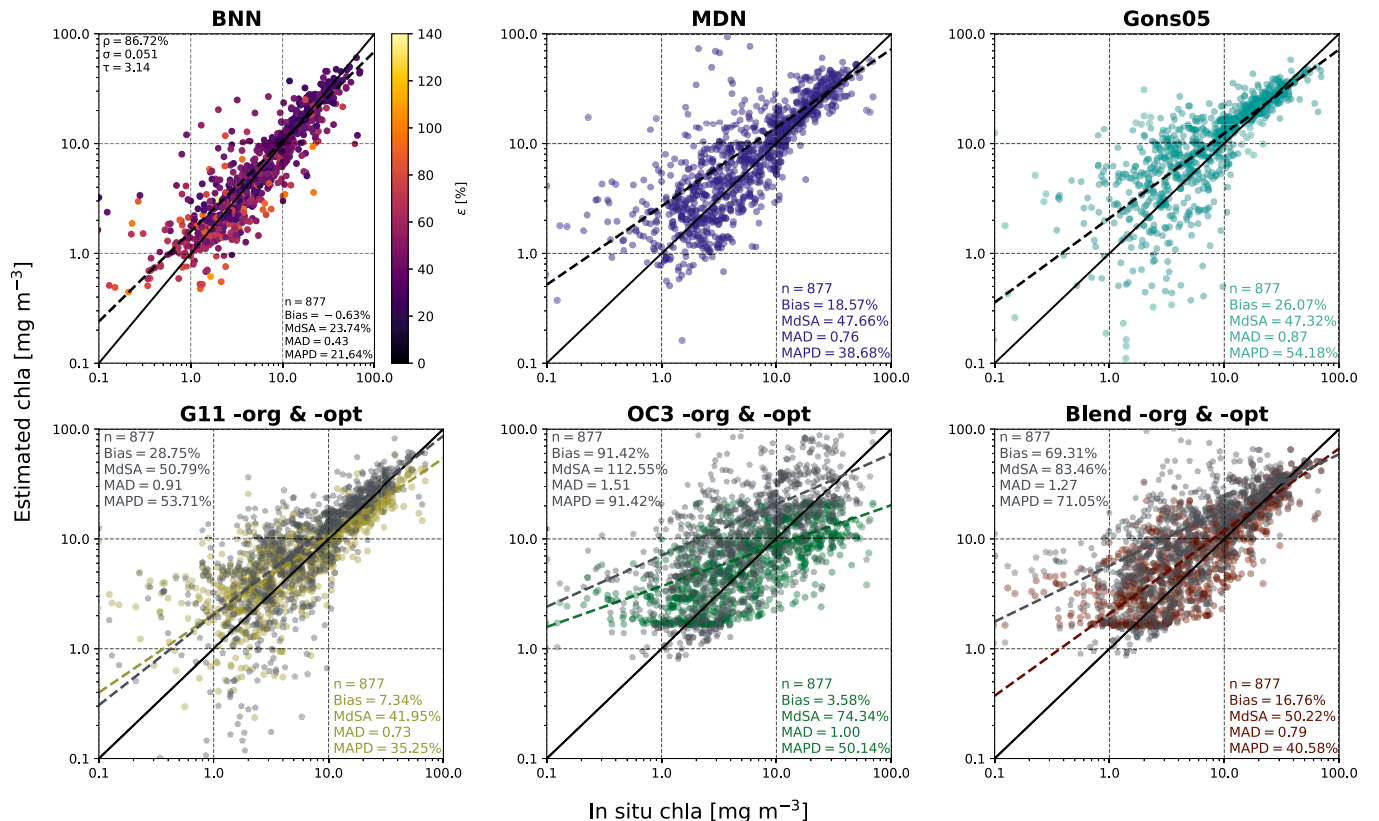


Fig. 5. Chl_a retrieval results by the OLCI BNN and reference algorithms in the 50/50 training/test split assessment. For G11, OC3 and Blend chl_a retrievals through original coefficients (-org) are shown in grey, and optimised coefficients (-opt) in colour.

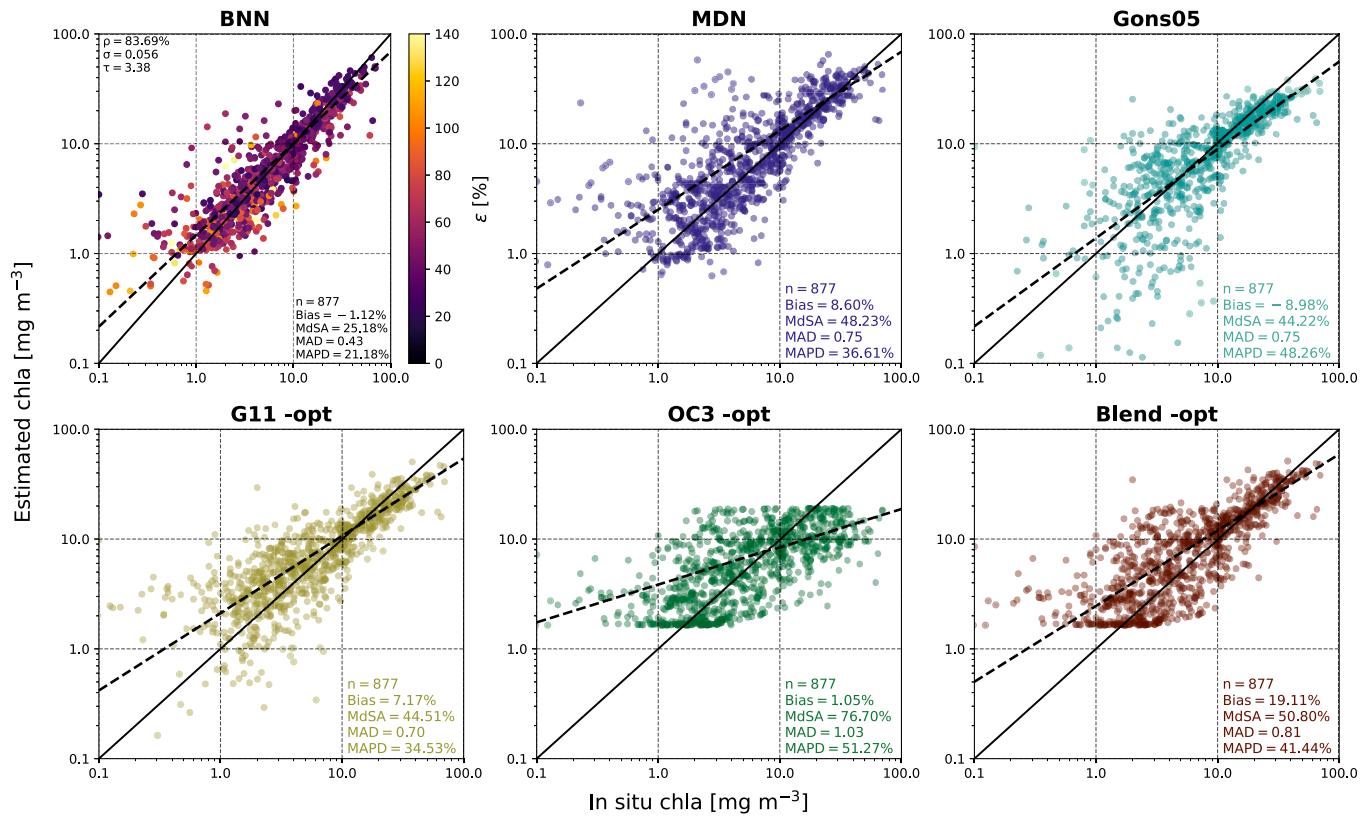


Fig. 6. Chla retrieval results by the MSI BNN and reference algorithms in the 50/50 training/test split assessment.

Table 1

Retrieval accuracies of the chla algorithms per TS class for the test set, compared through the MdSA (in %) metric. Lowest MdSA achieved by an algorithm for each TS in bold.

Algorithm	Oligotrophic (chla ≤ 8 [mg m ⁻³]; n = 548)	Mesotrophic (chla >8 and ≤ 25 [mg m ⁻³]; n = 253)	Eutrophic (chla >25 [mg m ⁻³]; n = 76)
	OLCI		
BNN	28.78	18.95	17.28
MDN	76.43	25.33	19.37
Gons05	109.33	30.00	13.98
G11 -org	110.35	35.08	17.40
G11 -opt	75.09	19.26	34.85
OC3 -org	154.43	80.95	57.31
OC3 -opt	78.95	44.97	191.88
Blend -org	147.98	49.73	16.88
Blend -opt	83.21	24.74	19.84
	MSI		
BNN	32.42	19.14	19.77
MDN	71.13	26.43	20.94
Gons05	81.57	20.28	46.41
G11 -opt	72.58	17.52	35.98
OC3 -opt	79.54	47.20	207.57
Blend -opt	81.97	27.15	21.27

calibration by the OLCI and MSI BNNs was accurate (MACD 0.051 and 0.056, respectively) and PICP included $>83\%$ of the *in situ* chla measurements. Sharpness of 3.14 (OLCI) and 3.38 (MSI) across the large concentration range was moderate - low. Highly uncertain chla estimates were identified by both BNN configurations.

4.2. Leave-one-out: BNN generalisation ability

In the LOO assessment, the OLCI BNN MdSA was expectedly higher for some regions (MdSA between 19 and 77%) than for the 50/50 training/test split assessment (Fig. 7). In comparison the OLCI BNN

outperformed the reference algorithms in 7/10 regions and remained competitive for the other regions. Highest gains in retrieval accuracy by the OLCI BNN were made for region 4 consisting of the UK, Sweden, Finland and Estonia (MdSA 19% lower), the South African reservoirs representing region 7 (MdSA 18% lower) and Lake Erie, Ontario, and Winnipeg of region 10 (MdSA 40% lower). For most regions, MSI MdSA was only slightly higher than for OLCI. However, the largest difference in retrieval accuracy between the MSI and OLCI BNN configurations was linked to regions 4 and 10, whereby the MSI configuration was about 20% less accurate. The results of the independent region evaluation emphasise that measurements included in this dataset are representative for other regions of the globe. Consequently the dataset enabled a robust

Table 2

Retrieval accuracies of the chla algorithms for OLCI (top) and MSI (bottom) for the test set per OWT class, compared through the MdSA (in %) metric. Lowest MdSA achieved by an algorithm for each OWT in bold.

Chla algorithm	OWT 2 (n = 261)	OWT 3 (n = 257)	OWT 4 (n = 149)	OWT 5 (n = 13)	OWT 9 (n = 197)
OLCI					
BNN	23.21	27.73	20.82	42.51	23.67
MDN	36.87	64.72	47.53	109.67	44.73
Gons05	37.48	136.90	31.90	60.51	49.07
G11 -org	37.98	148.50	42.58	67.88	48.68
G11 -opt	32.71	82.80	34.47	47.48	39.66
OC3 -org	146.81	74.03	134.55	262.57	135.89
OC3 -opt	65.25	66.49	134.44	443.42	57.21
Blend -org	118.24	64.83	66.59	44.39	96.49
Blend -opt	41.11	69.44	41.31	24.82	55.09
MSI					
BNN	24.27	23.14	18.22	19.67	28.49
MDN	30.16	67.32	51.76	83.30	41.10
Gons05	32.02	107.92	40.09	63.58	66.14
G11 -opt	29.72	63.43	39.23	36.52	46.04
OC3 -opt	72.57	59.26	135.18	433.24	57.65
Blend -opt	44.06	102.98	36.62	37.20	51.26

generalisation of the BNNs to systems that shared the same OWTs, but that were from regions not included in the respective training dataset.

Through LOO, the effectiveness of the uncertainty calibration can be analysed (Fig. 8). MACD and Sharpness did not co-vary linearly with chla accuracy. For example, OLCI MdSA in regions 2, 5 and 9 was the smallest, but the corresponding uncertainty metrics were not consistently small, too. In region 9, chla accuracy with MdSA <20% was achieved, but MACD (> 0.20) and Sharpness (> 5) were high in comparison to other regions. Nevertheless, region 9 PICP was 100%, indicating that all BNN chla estimates fell within the estimated confidence intervals. Overall, OLCI PICP was on average 7% higher than MSI (85.24% and 78.36%, respectively). Although the MSI BNN was only slightly less accurate in the retrieval of chla, the algorithm was less certain about its estimates resulting in a lower PICP percentage and higher MACD.

4.3. Lake Geneva match-ups

Satellite-derived $R_{rs}(\lambda)$ for *in situ* match-ups over Lake Geneva were generated through POLYMER and C2RCC applied to OLCI and MSI products. POLYMER calculated negative values in several red bands of the spectrum for 21 OLCI and 12 MSI match-ups. To assure a better match-up comparison of the chla algorithm performances, the 21 OLCI and 12 MSI observations obtained through POLYMER were excluded. Further, Gons05 and G11 were excluded because these red/NIR algorithms showed the largest retrieval uncertainties in the 50/50 split and LOO assessments for chla <10 mg m⁻³. Chla retrievals from OLCI POLYMER and C2RCC were more accurate than for MSI from the evaluated algorithms (see Table 3). OLCI POLYMER OC3 -org and Blend -org configurations were most accurate (MdSA 48.41% and 36.61%, respectively), closely followed by the BNN (MdSA 49.75%). Optimising the OC3 and Blend algorithms coefficients did not improve the chla estimation for the match-ups, unlike in the 50/50 split and LOO assessments. The optimisation of the OC3 and Blend algorithm coefficients systematically increased the underestimation of chla (Bias < -70%). The reason for this difference lays in the large value range used for the algorithm coefficient optimisation, which was naturally higher than the variation in Lake Geneva. Algorithms using OLCI imagery corrected through POLYMER underestimated chla (Bias between -21% for BNN and -175% for MDN). For OLCI C2RCC match-ups, the BNN was four times more accurate than the OC3 and Blend configurations (Fig. 9). The reference algorithms performed worse on OLCI observations compared to the *in situ* 50/50 training/test set split and LOO assessments, which indicates that the reference algorithms were more affected by AC uncertainties than the BNN. Contrary to POLYMER, C2RCC OLCI and MSI BNNs were more balanced (Bias -0.48% and 22.20%, respectively), resulting in less variability of estimated chla (Fig. 10). PICP was high (> 76%) across all four sensor and AC combinations, and MACD values were similar to the previous LOO analysis on independent region data. Sharpness values were lower than in the 50/50 split and LOO assessments, indicating small

average standard deviations and thus a better uncertainty calibration. OLCI and MSI MDNs systematically underestimated chla across all four sensor and AC configurations (MdSA >175%, Bias < -175%).

4.4. Lake Geneva time series

Phytoplankton vertical variability is seasonal in Lake Geneva (Min-audo et al., 2021) and a major factor driving optical properties (Nouchi et al., 2018). The impacts of phytoplankton vertical variability and its effect on remote sensing retrievals is still poorly understood. Seasonal variations over the low to moderate chla range (0–8 mg m⁻³) were accurately estimated by the OLCI BNNs during the four-year period, independent of the prior AC algorithm. The confidence intervals in most cases (PICP >80%) captured the corresponding *in situ* chla value (see Figs. 11, 12). For a spring bloom in 2020, moderate chla (8–12 mg m⁻³) was not estimated with similar precision when compared to low chla across the four years. The confidence intervals were too small and did not contain most of the high peaks during the bloom. Conversely, *in situ* estimates in August–September 2021 with chla >5 mg m⁻³ were closely matched by the OLCI and MSI BNNs.

The MSI BNNs showed a larger spread in the low to moderate chla range compared to the OLCI BNNs, with wider confidence intervals. Larger and less precise intervals (higher Sharpness) via MSI resemble the results of the 50/50 split and the match-up assessments. The uncertainty estimates by the OLCI BNNs were throughout the exercises more accurately calibrated. Across the time series, $R_{rs}(\lambda)$ derived through POLYMER were associated with higher BNN chla uncertainty than through C2RCC. Between April and May 2021, POLYMER estimated negative $R_{rs}(\lambda)$ values in red bands (not shown) which led to higher uncertainty in chla estimates compared to more stable and lower chla estimates following from C2RCC for the same period.

4.5. OLCI and MSI BNN chla and uncertainty products

The BNNs were applied to several OLCI and MSI images over lakes with varying optical properties to assess horizontal consistency and visualise chla estimates alongside retrieval uncertainties. Single-day products over Lake Geneva, the Southern Island in New Zealand (SNZ), Lake Turkana in Kenya and boreal areas of Sweden and Canada were processed. Fig. 13 illustrates OLCI BNN chla through prior C2RCC AC and associated uncertainty products over Lake Geneva from March 2020, November 2020, and August 2021. The maps using the BNN match the time series chla patterns and ranges for Lake Geneva from the oligotrophic states in March 2020 to mesotrophic levels in August 2021 (Soullignac et al., 2018). The retrieval uncertainties ranged between 20 and 60% in Lake Geneva similar to the results of the previous match-up and time series assessments. The results over the LÉXPLORE location can thus be considered transferable to other regions of the lake. During March 2020, areas with high uncertainty are observed near the north

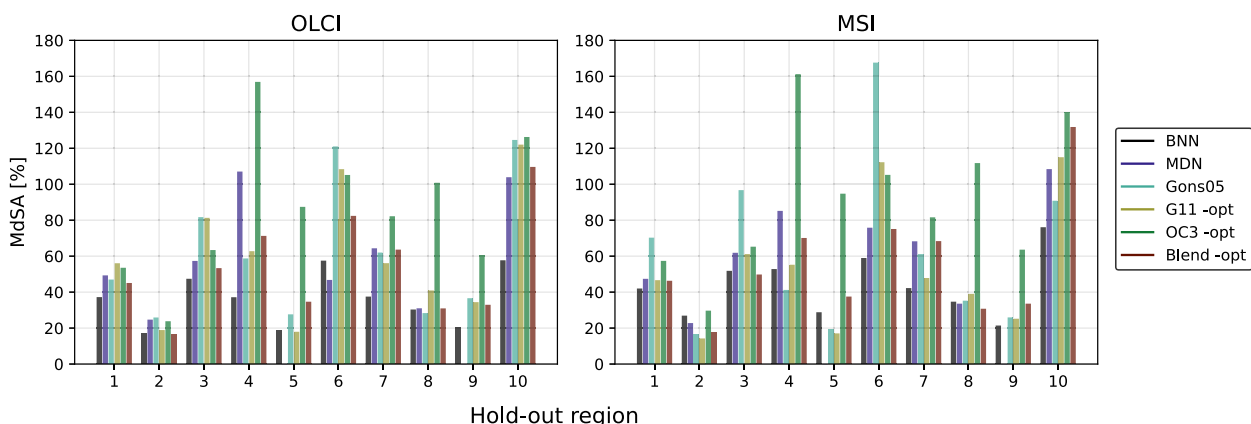


Fig. 7. Change in MdSA of the chla algorithms for each region as part of the LOO assessment.

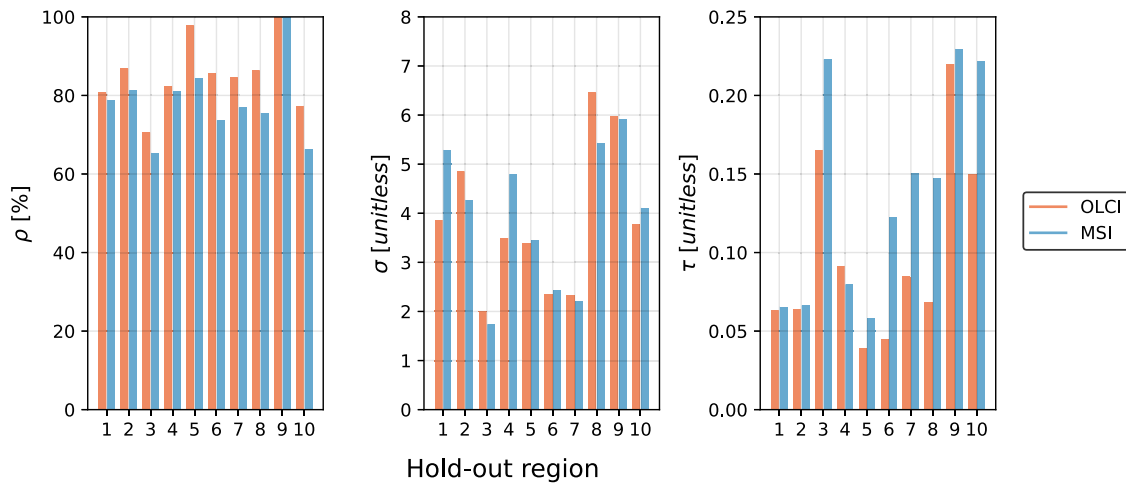


Fig. 8. BNN uncertainty calibration metrics PICP (ρ), Sharpness (σ) and MACD (τ) per region as part of the LOO assessment.

Table 3

Performance metrics of the chl_a algorithms for the OLCI and MSI match-ups over LÉXPLORE. Results for POLYMER are coloured in orange, C2RCC in blue. Lowest MdSA per combination in bold.

OLCI POLYMER / C2RCC

Algorithm	Bias [%]		MdSA [%]		MAD []		MAPD [%]	
BNN	-21.16	-0.48	49.79	59.07	0.59	0.71	36.51	44.53
MDN	-175.18	-318.68	175.18	323.23	1.76	3.39	63.66	76.51
OC3 -org	-44.24	-253.43	48.41	253.43	0.53	2.43	34.33	67.70
OC3 -opt	-76.27	-193.06	76.27	193.06	0.83	2.70	43.27	65.88
Blend -org	-23.75	-202.80	36.61	209.60	0.54	2.29	28.23	67.70
Blend -opt	-73.40	-226.33	78.22	235.89	0.90	2.95	44.08	70.23

MSI POLYMER / C2RCC

Algorithm	Bias [%]		MdSA [%]		MAD []		MAPD [%]	
BNN	-84.26	22.20	84.26	60.72	0.75	0.84	45.73	51.48
MDN	-283.02	-634.84	283.02	634.84	2.50	7.60	73.89	87.30
OC3 -org	-87.66	-535.99	93.65	541.88	1.03	8.35	49.61	84.42
OC3 -opt	-137.40	-370.75	137.40	370.75	1.30	5.14	57.88	78.87
Blend -org	-213.79	-1471.02	213.79	1471.02	2.65	27.38	71.87	94.20
Blend -opt	-133.36	-373.17	133.36	381.39	1.24	8.54	57.88	79.23

shore of Lake Geneva, likely associated with contamination of water-leaving radiance by bottom reflectance, adjacent land, or both.

OLCI products processed with C2RCC and POLYMER for SNZ are shown in Fig. 14. BNN chl_a products varied spatially between the lakes and AC algorithms. Moreover, uncertainties of C2RCC derived-chl_a were significantly higher (between 50 and 100%) than for POLYMER (Fig. 14). Unlike for Lake Geneva, POLYMER did not produce negative $R_{rs}(\lambda)$ in any wavebands over this area. A surface measurement of *in situ* chl_a in Lake Hawea (stars on Fig. 14) was 0.89 mg m^{-3} on the 9th of February during the Sentinel-3 overpass, which closely matched the BNN chl_a of this scene. Overall, highest uncertainties were generated for chl_a $< 1 \text{ mg m}^{-3}$, a range for which training data were scarce. The scarcity became apparent in the top left corner of the imagery in the clear coastal waters. The uncertainties for chl_a $< 1 \text{ mg m}^{-3}$ were similar to the results of the initial 50/50 split analysis. Highest chl_a retrieval uncertainties from the BNN and the reference algorithms are observed in this concentration range.

The MSI BNN generated high-quality chl_a products for small water bodies (SWBs). Fig. 15 contains MSI BNN products over SWBs of Canada and Sweden. Satellite-derived $R_{rs}(\lambda)$ from SWBs are prone to adjacency effects which degrade the quality of the retrievals (Paulino et al., 2022; Sterckx et al., 2011). In Fig. 15 D-F, small lakes contained water pixels (white squares) for which the MSI BNN produced high uncertainties. Associated reflectance spectra of small lakes in this scene used to retrieve chl_a were influenced by the proximity to land, resulting in high uncertainty. The MSI BNN products demonstrate sensitivity to changes in chl_a between the lakes captured on this scene. Consistent with OLCI BNN retrievals, lake areas depicted in MSI BNN products exhibited higher uncertainties where chl_a was lowest (Fig. 15 A, B, C). In November 2021 (Fig. 15 G-I) the RGB of “Clearwater Lake” appears extremely bright (see yellow square), which the BNN associated with high chl_a uncertainty. OLCI BNN products from December 2021 and January 2022 (Fig. 16) cover the largest desert lake in the world, Lake

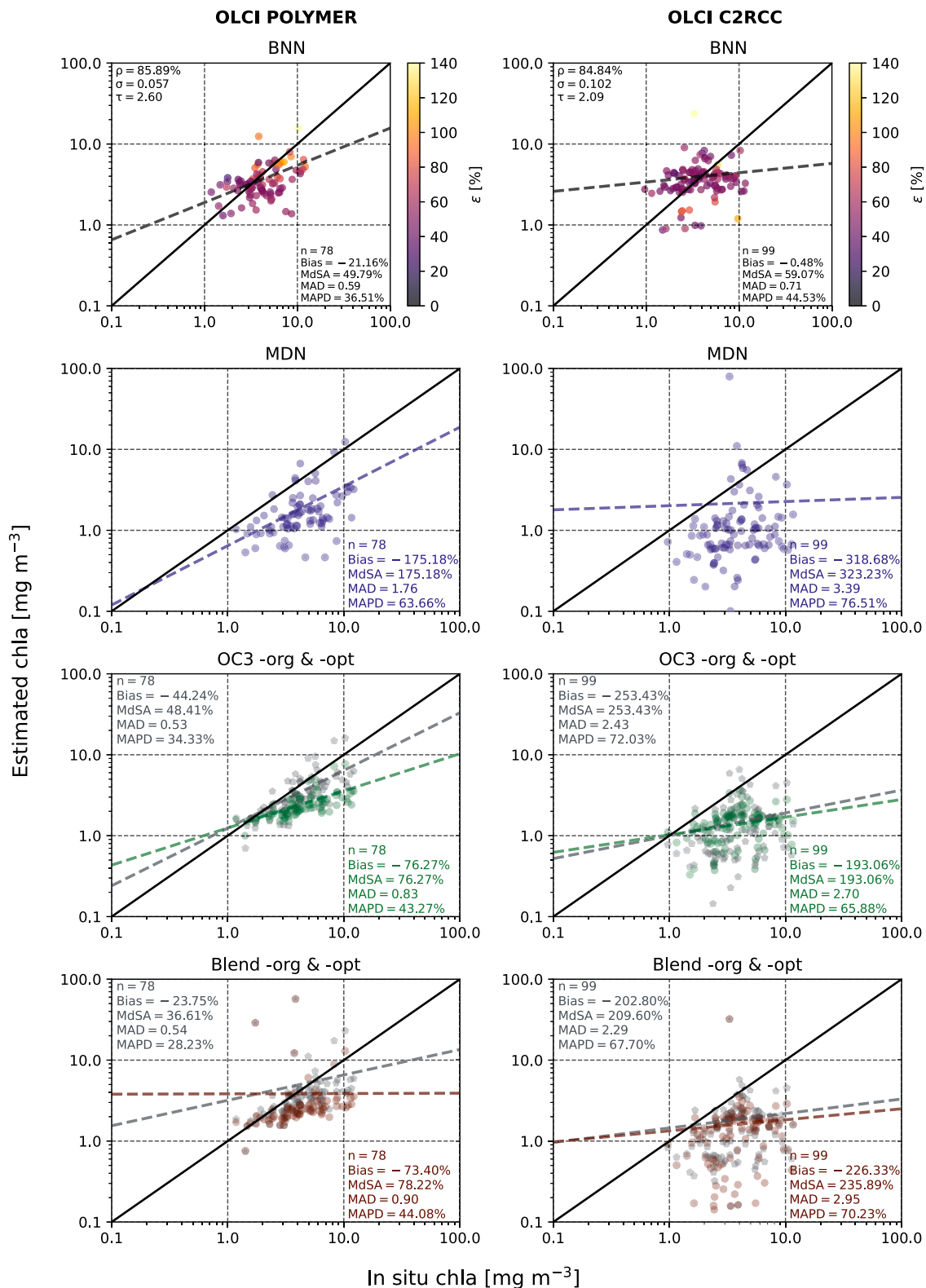


Fig. 9. LÉXPLORE platform (Lake Geneva) OLCI match-up chl a estimates from BNN, MDN, OC3 -org, OC3 -opt, Blend -org and Blend -opt algorithms. For OC3 and Blend, grey markers indicate the original and coloured markers the optimised algorithm coefficients.

Turkana (Kenya), known for its spatially variable optical properties (Liu et al., 2021; Tebbs et al., 2020). OLCI BNN products over Lake Turkana documented a gradient in chl a of the more turbid waters in the north of the lake that decreases towards the lake's southern end. BNN chl a following C2RCC (Fig. 16 B, I) was higher ($2\text{--}6 \text{ mg m}^{-3}$) than through

POLYMER ($2\text{--}3.5 \text{ mg m}^{-3}$; Fig. 16 E, L). Chl a uncertainties from C2RCC were approximately 50% lower than by POLYMER for both days.

The product uncertainties were used to filter retrievals and exclude areas of water bodies for which the BNN assigned high uncertainty. $R_{rs}(\lambda)$ from POLYMER caused the BNN to produce unrealistic patches of

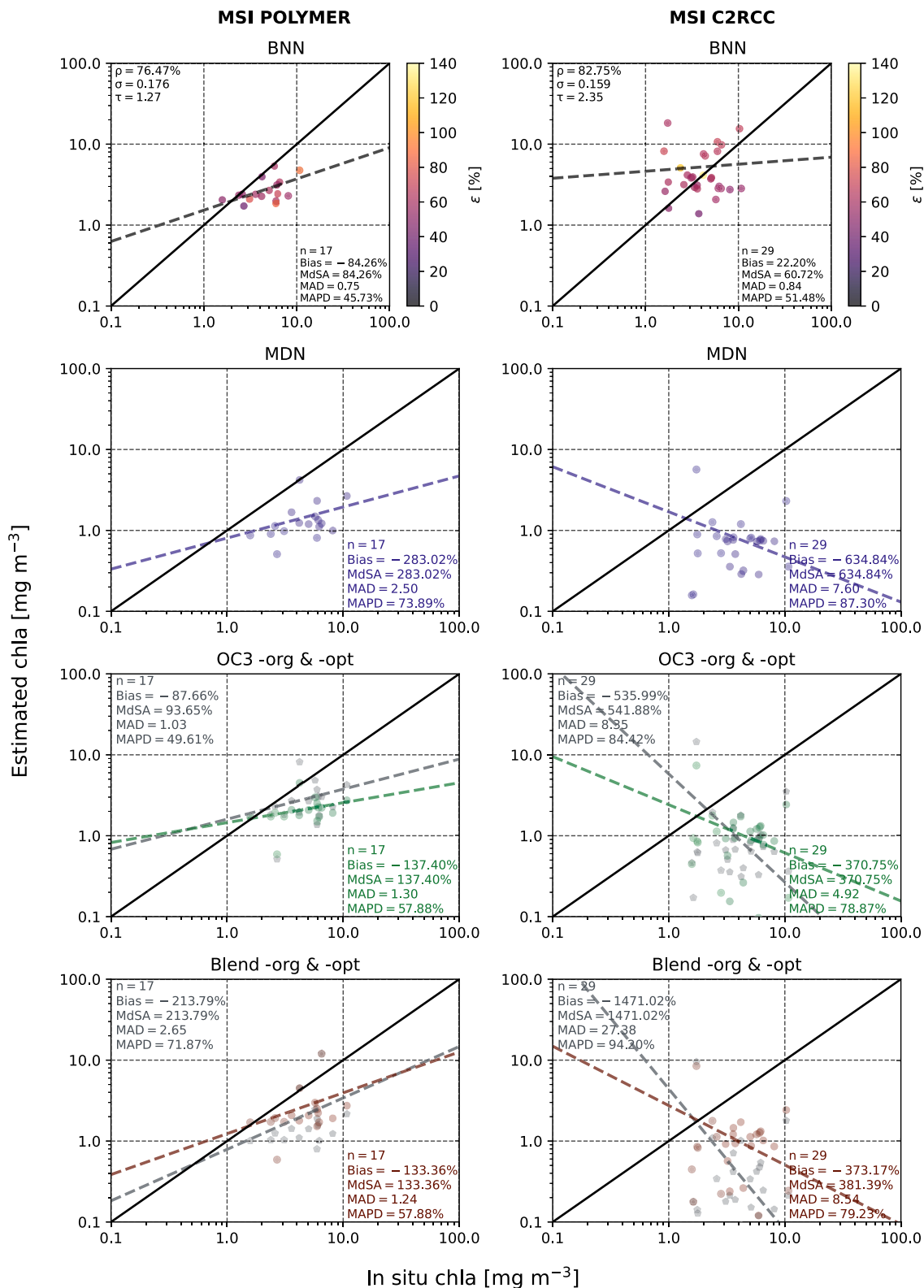


Fig. 10. LÉXPLORE platform (Lake Geneva) MSI match-up chl a estimates from the BNN, MDN, OC3 -org & -opt and Blend -org and -opt algorithms. For OC3 and Blend, grey markers indicate the original and coloured markers the optimised algorithm coefficients.

chl a around 1 mg m^{-3} (Fig. 16 F, M). These patches did not resemble the optical gradient apparent in the area and were not produced through prior C2RCC AC. In contrast, for the Jan. 2022 product C2RCC reflectance spectra caused the BNN to generate unrealistically low chl a in the middle of the lake and nearshore when compared to neighbouring pixels

(Fig. 16 I). The unusual BNN chl a patches and areas caused by the POLYMER and C2RCC ACs were consistently associated with 40–100% higher uncertainties than surrounding pixels (Fig. 16 F, J, M). These highly uncertain areas over Lake Turkana were treated and resulted in new products. First, the corresponding pixels were removed by imposing

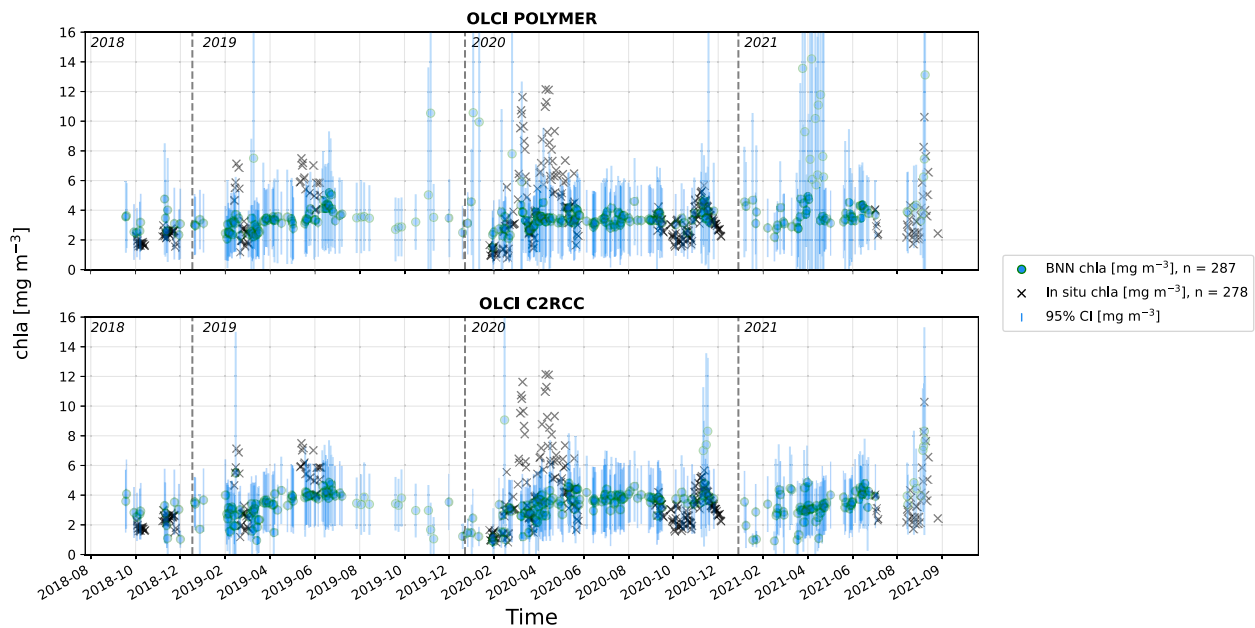


Fig. 11. BNN OLCI time series of LéXPLORE (Lake Geneva), including all match-ups.

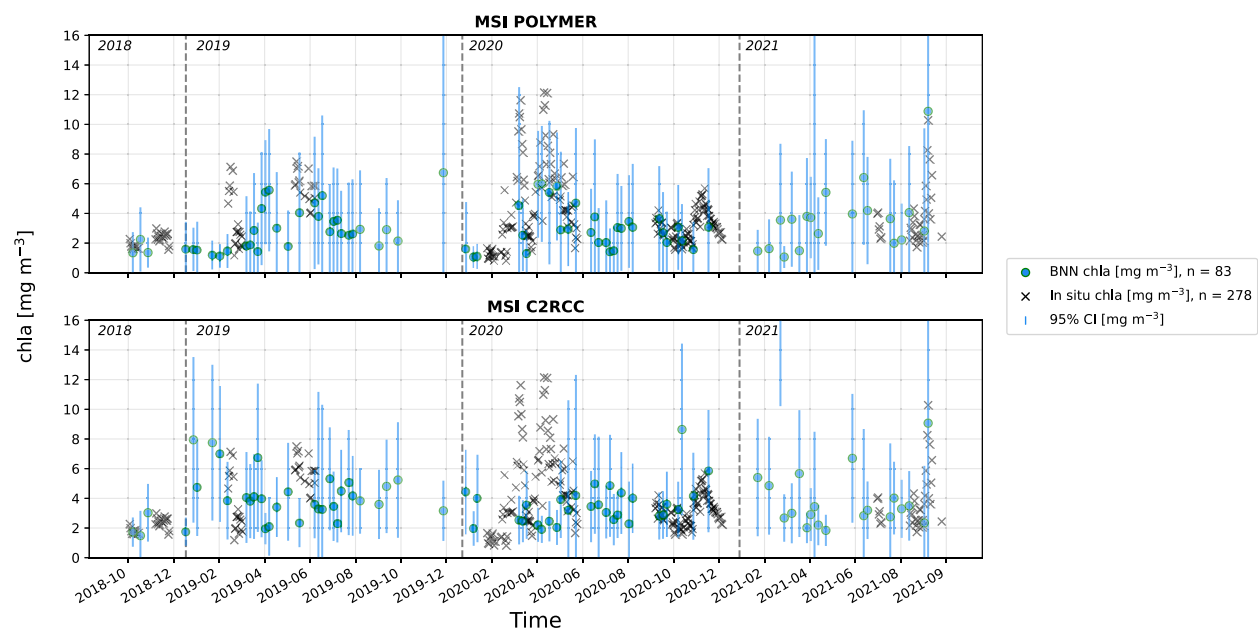


Fig. 12. BNN MSI time series of LéXPLORE (Lake Geneva), including all match-ups.

BNN uncertainty thresholds (higher than 85% for Dec. 2021 and 65% for Jan. 2022) on the entire scene. The procedure eliminated almost all observations with high uncertainty and resulted in more homogenous products (Fig. 16 D, G, K, N). Second, the BNN chl a uncertainties associated with each AC algorithm were used to generate merged chl a and uncertainty products (Fig. 17), by selecting the chl a result per observation with the lowest BNN uncertainty from either AC. The resulting, merged product replaced the highly uncertain areas of the AC processors (see Fig. 16 for the products used in the merge) and led to an improved product quality both in terms of chl a and uncertainty.

5. Discussion

Chl a estimation uncertainties have repeatedly been shown to be highest in oligotrophic and mesotrophic lakes (Liu et al., 2021; Neil et al., 2019; Werther et al., 2021). The retrieval results over the chl a

range considered here confirmed these findings. Uncertainty is common in lake remote sensing and unlike the reference chl a algorithms, the BNNs indicated when retrieved chl a was uncertain. Here we discuss the BNN chl a and uncertainty estimation, the performance of the reference algorithms, the uncertainty calibration and the handling of uncertain estimates as exemplified for Lake Turkana.

5.1. OLCI and MSI BNN chl a estimation

The uncertainty associated with BNN chl a varied between the five performance assessments. Lowest BNN chl a retrieval uncertainty was found for the 50/50 split assessment over established chl a algorithms. The 50/50 split performance of the BNN resembles results reported in the MDN publications (O'Shea et al., 2021; Pahlevan et al., 2020). In this assessment strategy, the systems in both training and test sets are not strictly independent. Observations from the same system can occur in

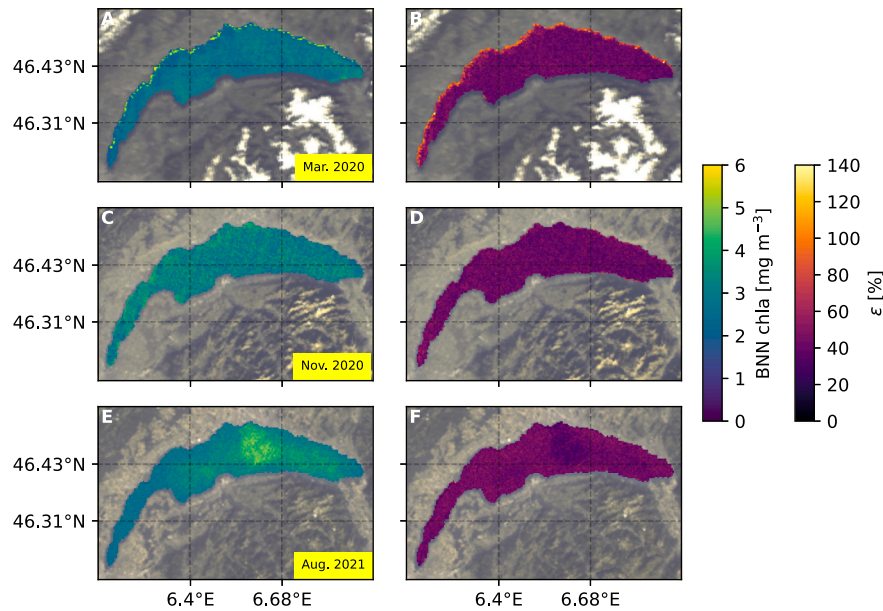


Fig. 13. OLCI BNN chl a and uncertainty products obtained through C2RCC AC over Lake Geneva on 13th of March 2020 (A, B), 11th of November 2020 (C, D) and 21st of August 2021 (E, F).

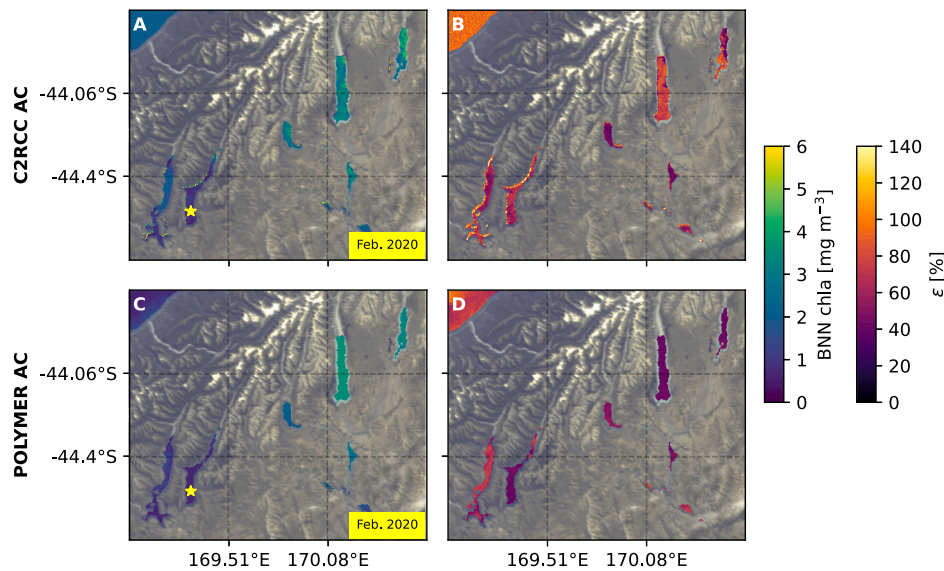


Fig. 14. OLCI BNN chl a and uncertainty products over southern New Zealand, 9th of February 2020. Chl a was measured *in situ* at the location of Lake Hawea (star symbol) as 0.89 mg m^{-3} during the overpass.

both sets due to the prior randomisation of the entire dataset, thereby introducing knowledge about similar optical properties in the algorithm training stage. Controlled randomisation prevents dataset shifts and the 50/50 dataset split results therefore represent an idealistic evaluation scenario. However, dataset shifts are common during satellite application of a developed ML algorithm. Therefore, the region-wise LOO assessment was conducted, and the results provided further insight into the BNN capabilities. The LOO assessment corroborated the 50/50 split results, whereby for 7 out of 10 regions the BNN outperformed reference algorithms. Expectedly, overall accuracy for LOO was worse than in the 50/50 split assessment. Regional chl a accuracy varied from 19% to nearly 77% MdSA for the BNNs but was consistently higher for the reference chl a algorithms. Moreover, the reference chl a algorithm performances varied strongly between regions. LOO results highlight that the training dataset of this study is representative for measurements sharing the same OWTs, even if from different regions. The LOO results

also suggest that varying degrees of uncertainty between the observations of the dataset exist that propagated into the BNN chl a estimation. This source of observational uncertainty is irreducible for the BNNs and sets a realistic baseline for expected performance. Compared to the *in situ* assessments, overall BNN accuracy was expectedly lower for the match-ups over Lake Geneva. Although there has been a decline in performance, this was not surprising given that the ACs induced uncertainty ranging from 18 to 60% in the blue to red bands (see Appendix 4). The MdSA of the OLCI BNN match-ups over Geneva were 49.7% (POLYMER) and 59.07% (C2RCC) which is considered within an acceptable performance range. The decrease in accuracy was greater for MSI because the BNN model was overall less precise, as was also seen in the *in situ* data analyses. Additionally, this is consistent with a decline in the accuracy of MSI AC-derived R_{rs} when compared to OLCI.

BNN chl a uncertainties were highest for chl a $< 1 \text{ mg m}^{-3}$ across larger and smaller water bodies captured on OLCI and MSI products over

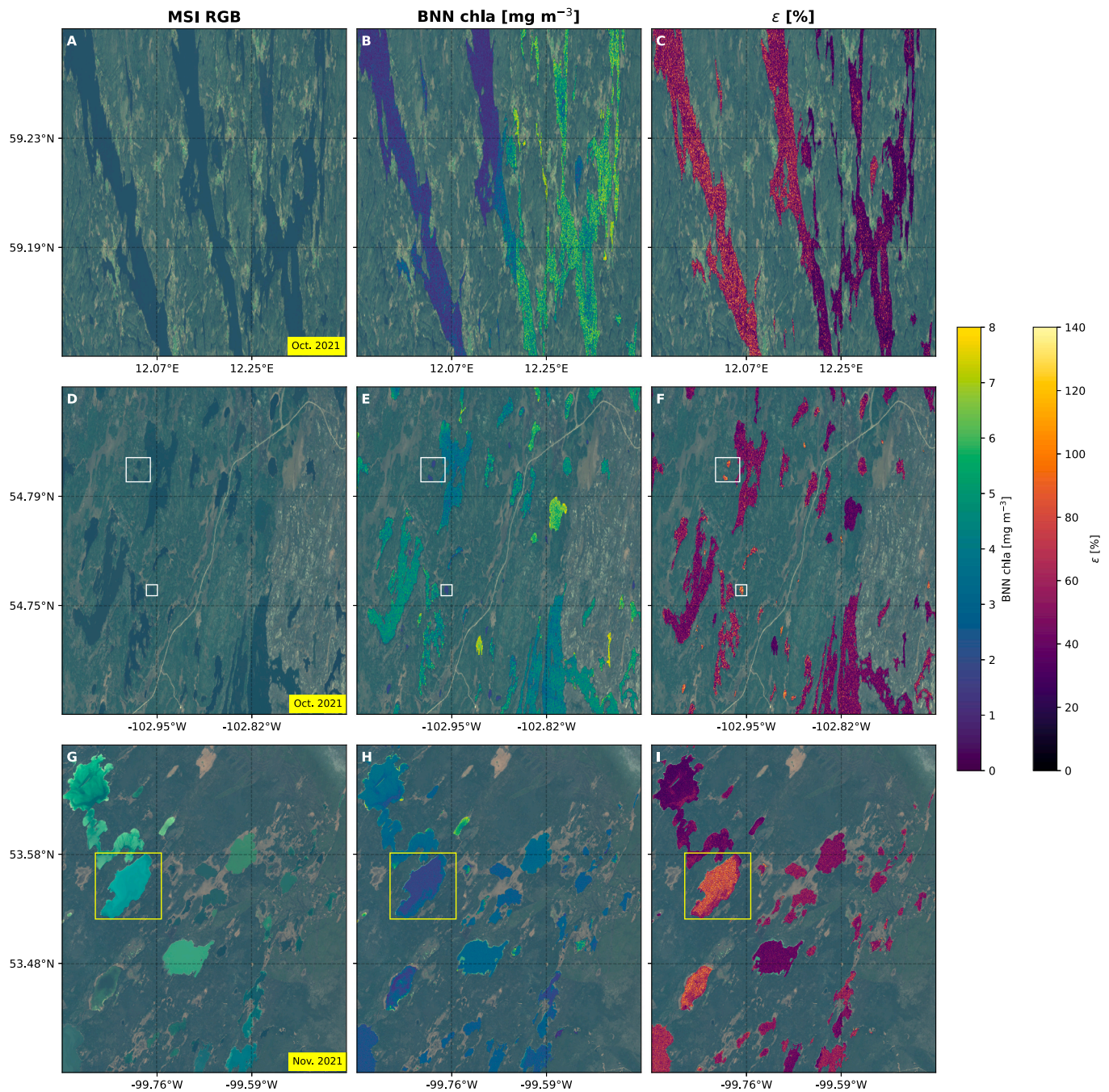


Fig. 15. BNN-C2RCC chl a and uncertainty products for MSI (20 m) over Sweden (A, B, C) on 11th of October 2021, and Canada on 16th of October (D, E, F) and 4th of November 2021 (G, H, I). A, D, G represent MSI L1 products with true water colours obtained from top-of-atmosphere sensor radiance, while B, E, H depict produced chl a. C, F, I represent associated chl a uncertainty. White and yellow boxes surround small water bodies with high BNN uncertainty. See explanations in this section for the boxes.

SNZ, Canada and Lake Turkana. Because measurement availability is globally scarce in the clearest lakes, measurement coverage was low in the training dataset. The low amount of training data and resulting high chl a uncertainty over this range clearly indicate a measurement gap. Radiative transfer simulations or coastal data sets may be able to fill this measurement gap to further reduce chl a estimation uncertainties. In addition to measurement scarcity, systematic measurement error is prone to lowest chl a and phytoplankton absorption coefficients (McKee et al., 2014). With regards to higher biomass conditions, only 9.5% of the dataset consisted of chl a $>25 \text{ mg m}^{-3}$ (maximum 68 mg m^{-3}), therefore chl a estimation uncertainty will rise with increasing chl a. Oligo- and mesotrophic waters most likely do not exceed this threshold

even when experiencing seasonal, short-term phytoplankton blooms such as those observed in Lake Geneva. Observations characterised by high CDOM ($> 2 \text{ m}^{-1}$) were limited in the collated dataset. High CDOM is common in boreal and tundra lakes with low - moderate chl a but these systems are globally under-sampled and as such only constitute a small part of the training data available for the BNNs.

5.2. Performances of the chl a algorithms

We compared the constructed BNNs to five reference chl a algorithms that differed in their approach to estimating chl a. The OLCI and MSI MDNs were created using a bigger community-dataset ($n = 2943$) than

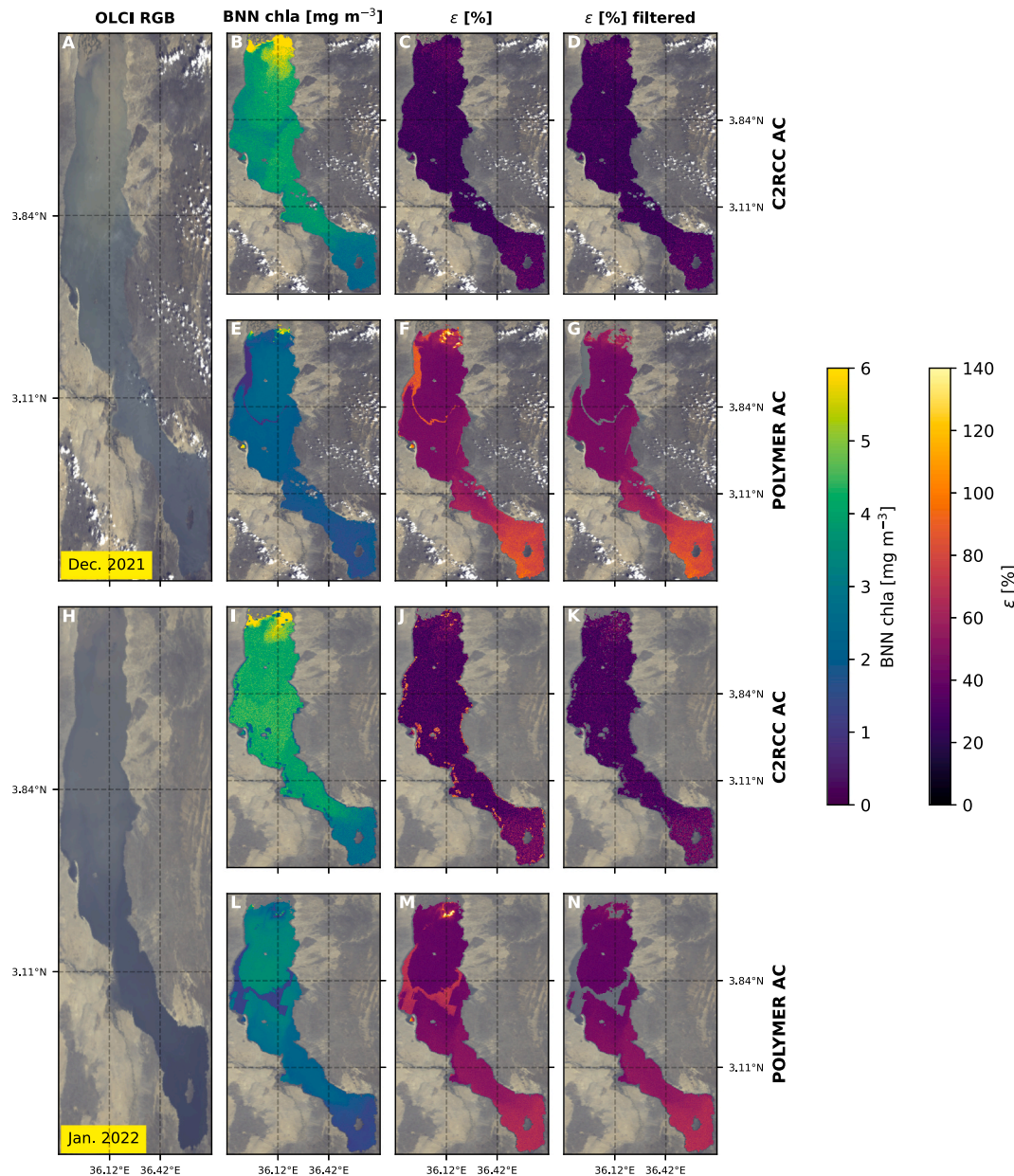


Fig. 16. OLCI BNN chl a and uncertainty products for Lake Turkana (Kenya) on 13th of December 2021 and 14th of January 2022 using C2RCC and POLYMER ACs. A, H: OLCI L1 RGB. B, E, I, L: BNN chl a. C, F, J, M: Associated uncertainty. D, G, K, N: Filtered pixels based on uncertainty higher than 85% for Dec. 2021 and 65% for Jan. 2022.

the BNNs ($n = 1755$). In contrast to the dataset used in this work, the MDN training dataset was centred on eutrophic waters (chl a dataset mean 21.7 mg m^{-3} , median 8.9 mg m^{-3} and standard deviation of 47.5 mg m^{-3}). Accordingly, the performance of the MDNs became steadily more accurate with increasing chl a for the observations of this study. A recent chl a retrieval uncertainty characterisation of the MDNs by Werther et al. (2022) revealed that MDN overestimation in low biomass waters (chl a $0\text{--}10 \text{ mg m}^{-3}$) is primarily caused by a lack of training observations from waters with moderate - high NAP and CDOM absorption.

To achieve higher generalisation across global lakes, it appears that training sets for ML algorithms must be further expanded, however gathering *in situ* measurements will remain expensive and time-consuming. As a result, novel techniques to overcome the lack of *in situ* data are required. Recent developments grounded in semi-supervised (Zhu and Goldberg, 2009) or self-supervised learning (Jing and Tian, 2021) extend the training set size through the transformation of unlabelled observations into available training data. Both learning approaches show promise in adjacent

remote sensing fields (Wang et al., 2022), but their potential remains untapped for aquatic remote sensing of biogeochemical proxies.

The red/NIR algorithms using 665 and 708 nm such as G11 -opt, Gons05 and Blend became highly accurate ($\text{MdSA} < 20\%$) at chl a $> 10 \text{ mg m}^{-3}$. The results match the findings of earlier studies across larger chl a ranges reporting inaccuracies of red/NIR algorithms in oligotrophic concentration ranges (Neil et al., 2019; Pahlevan et al., 2020; Werther et al., 2022). The OC3 -org and -opt versions showed moderate - high estimation uncertainties ($\text{MdSA} > 44\%$) across all trophic states in the 50/50, LOO and match-up assessments. For oligo- and mesotrophic lakes with highly varying OAC concentrations, phytoplankton estimation based only on a blue/green band ratio appears severely constrained.

The developed BNNs are an ensemble of individual NNs, with the size of the ensemble specified by the parameter S . Through MC dropout, S network variants (50 in this study) with differing numbers of neurons were produced, resulting in a varied chl a estimate for an identical input observation. Ensembles consistently outperform individual algorithms, either

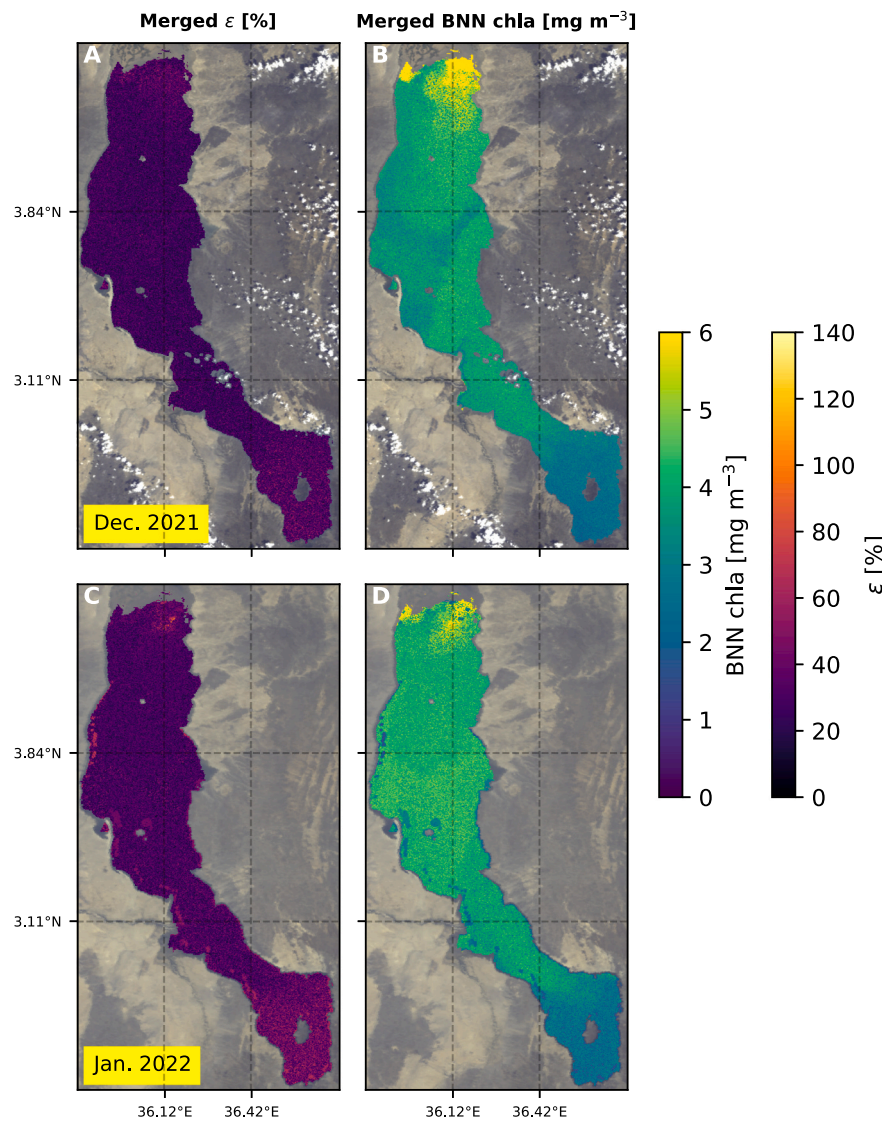


Fig. 17. OLCI BNN merged uncertainty and chl-a products from 13th of December 2021 and 14th of January 2022 over Lake Turkana. See Fig. 16 for the products used for the merge.

when combined through OWTs (Neil et al., 2019; Liu et al., 2021), statistical merging or blending (Smith et al., 2018; Schaeffer et al., 2022) or stacking of individual algorithms (Werther et al., 2021). However, using multiple algorithms increases model complexity at the expense of feature interpretability. The 25% dropout probability determined S distinct NN variants in this study. The 50 distinct NNs each had a different number of neurons (because different neurons were dropped in each NN variant). Due to their uniqueness, the importance of individual features (the multispectral bands) varied among the 50 NNs. The final chl-a value was determined by the chl-a distribution created by all of the different NNs. As a result, this chl-a distribution is a blend of distinct feature importances, making it a complex task to separate the individual NN contributions in order to identify which bands contributed meaningfully to the model.

5.3. Calibration of uncertainty

The quality of the uncertainty calibration was measured through the PICP, Sharpness and MACD metrics. Across the different assessments, the confidence intervals covered the *in situ* reference chl-a value in 3/4 cases (PICP >75%). However, the provided uncertainty estimate was not consistently well-calibrated across the assessments. The Sharpness metric specifically measured the range of the uncertainty estimate per observation

and small Sharpness values were not achieved in all performance assessments. Moreover, even though OLCI performance did not differ significantly from MSI, the OLCI uncertainties were consistently better calibrated. More research is needed to understand under which conditions the BNN uncertainty calibration was systematically inaccurate and how it correlated with optical properties and training data availability.

The source of uncertainty can be the underlying BNN statistical model (the epistemic uncertainty) or an external source (the aleatoric uncertainty), such as prior AC, an adjacency effect or other unknown error as part of the measurements contained in the dataset (IOCCG, 2019). Whereas epistemic uncertainty can be improved through better data coverage and incremental improvements to the underlying statistical model, aleatoric uncertainty sources are often more difficult to identify, quantify and to correct for as they are methodologically de-coupled from the BNN. AC, for example, is often performed using a distinct algorithm. Uncertainty in the AC method for calculating $R_{rs}(\lambda)$ may thus propagate into the BNN chl-a estimation. It is difficult to identify and correct for this source of aleatoric uncertainty impacting BNN chl-a estimation accuracy without particular knowledge about the uncertainty introduced by the AC. To provide precise pathways for BNN improvements and to simplify the identification of uncertainty drivers, it is preferable to separate the aleatoric and epistemic uncertainty contributions entirely. Kendall and

Gal (2017) conducted preliminary work in this direction using MC dropout in computer vision tasks, but it is still far from established research practise in the field of remote sensing.

Besides the investigation into uncertainty origins, computational methods exist to improve the uncertainty calibration as a post-processing step. The approaches usually require a separate dataset, which in practise can be a subset of the overall data that was set aside, and which is consequently excluded from training or evaluation. Example techniques include histogram binning (Zadrozny and Elkan, 2001), isotonic regression (Kuleshov et al., 2018; Zadrozny and Elkan, 2002) and Platt scaling (Platt, 1999).

5.4. AC algorithm selection through BNN uncertainty

Atmospheric correction is the greatest source of uncertainty in satellite remote sensing of OACs in lakes (Pahlevan et al., 2021a; Pereira-Sandoval et al., 2019; Wang et al., 2019; Warren et al., 2019). Observations associated with high uncertainty differed between the BNN estimates for the same products over Lake Geneva, SNZ and Lake Turkana. To deal with varying degrees of uncertainty introduced by AC algorithms, the BNN chl_a uncertainties were used as a selection mechanism by imposing thresholds (as undertaken in Fig. 16) or by retaining the observation associated with the lowest BNN uncertainty (Fig. 17). The former approach omits observations and is desirable when AC quality is unknown, during unsupervised processing of imagery or when a comparison to another AC algorithm is not possible. When two or more ACs for the same product are available, the latter approach to obtain a merged product is the favoured option. Moreover, the merging of products exemplified in Fig. 17 provides a pathway to decide between different ACs. Since AC performance largely varies with faced atmospheric properties and water column IOPs, it is *a priori* impossible to know which AC performs most accurate. A more thorough analysis with multiple products, areas and atmospheric conditions should be undertaken to determine whether AC selection through BNN uncertainty consistently improves the chl_a product quality.

6. Conclusion

The BNNs developed in this study represent a novel approach to derive chl_a and associated uncertainty estimates from $R_{rs}(\lambda)$ from Sentinel-2 (MSI) and Sentinel-3 (OLCI) over oligo- and mesotrophic lakes with the ability to account for phytoplankton abundance shifts into the eutrophic regime. Using the largest *in situ* dataset over oligo- and mesotrophic lakes and reservoirs available to date, the BNNs were trained, validated, and compared against state-of-the-art and community-established algorithms. In this study we demonstrated that the developed BNN algorithms work similarly for both MSI and OLCI observations (Bias <1.5%, MdSA <26%, MAPD <22%), suggesting their potential to produce high-quality chl_a products across multiple missions. We found that chl_a uncertainties for MSI were on average 10–25% higher than for OLCI reaffirming that additional bands on OLCI not only benefit the chl_a retrieval accuracy, but effectively reduce the associated uncertainty. Lowering uncertainty by increasing the number of available bands may be important for the growing number of hyperspectral Earth observation missions, such as the Italian PRISMA (PRecursore IperSpettrale della Missione Applicativa) and U.S. PACE (Plankton, Aerosol, Cloud and ocean Ecosystem).

As evidenced through long-term match-ups over Lake Geneva, the BNNs were able to improve upon the chl_a retrieval of established methods. Across all assessment strategies, the largest (MdSA > 25%) performance improvements by the BNNs were made over oligotrophic lakes (chl_a ≤ 8 mg m⁻³), whereas for chl_a > 25 mg m⁻³ established algorithm approaches performed similarly. As for any other chl_a algorithm relying on satellite-

derived $R_{rs}(\lambda)$, the chl_a accuracy of the OLCI and MSI BNNs varied with the prior AC algorithm. The BNN uncertainty estimate enabled to handle high uncertainty introduced by AC algorithms and other sources prior to chl_a retrieval. We exemplified common issues that are caused by ACs over oligo- and mesotrophic lakes: negative $R_{rs}(\lambda)$ in red bands, uncertainty induced by the adjacency effect and implausible $R_{rs}(\lambda)$ values stemming from extreme variation in bio-optical lake properties. To deal with inconsistent results obtained by the different AC algorithms, we suggest further research into the selection of AC products through BNN uncertainties to retain accurate chl_a retrievals. Moreover, research into strategies to improve the uncertainty calibration is recommended. Further downstream products depending on accurate chl_a, such as trophic status and net primary production, may impose uncertainty quality thresholds to retain high-quality chl_a retrievals. The BNN architecture presented in this study can be extended to other multi- and hyperspectral missions such as the Landsat missions, PRISMA and PACE and it supports the retrieval of other OACs and IOPs such as TSM and CDOM.

Code availability

The OLCI and MSI BNNs are available under https://github.com/mowerther/BNN_2022

CRediT author statement

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 Data curation: MW, ES, DO, ML, DG, TK, PH, AT, RG, AV
 Methodology: MW
 Formal analysis: MW
 Investigation: MW, ES, DO, SGHS, DG
 Software: MW
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Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Data availability

Data will be made available on request.

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Appendix 1. Dataset

Regions and inland water bodies constituting the dataset of this study. Colour shading matches Fig. 1.

Region	Countries	Dataset measurements (n = 1755)	Inland water bodies (n = 178)	References
1	Spain	109	Aguilar reservoir, Alarcon reservoir, Lake Albufera, Alcantara reservoir, Alcorlo reservoir, Almendra reservoir, Bornos reservoir, Burguillo reservoir, Canelles reservoir, Cernadilla reservoir, Cijara reservoir, Contreras reservoir, Cortes de Pallas reservoir, Cuerda del Pozo reservoir, El Atazar reservoir, Giribaile reservoir, Guadalcacin reservoir, Guadalen reservoir, Guadalteba reservoirs, Iznajar reservoir, La-Serena reservoir, Negratin reservoir, Pinilla reservoir, Rialb reservoir, Riano reservoir, Ricobayo reservoir, San Juan reservoir, Lake Sanabria, Santa Teresa reservoir, Terradets reservoir, Pantano de Tremp reservoir, Uribarri Ganboako urtegia reservoir, Valdecanas reservoir, Valmayor reservoir, Valparaiso reservoir, Valungo reservoir	(Ruiz-Verdú et al., 2005, 2008; Simis et al., 2007)
2	The Netherlands, Israel	107	IJsselmeer, Lake Kinneret	(Simis et al., 2005; 2007; Yacobi et al., 2011)
3	Switzerland, Italy	466	Lake Biel, Lake Garda, Lake Idro, Lake Como, Lake Iseo, Lake Geneva, Lake Maggiore, Lake Trasimeno	(Bresciani et al., 2011; Giardino et al., 2005, 2013, 2014, 2015; Guanter et al., 2010; Manzo et al., 2015)
4	United Kingdom, Sweden, Finland, Estonia	75	Lake Antu, Lake Bassenthwaite, Lake Coniston, Loch Katrine, Loch Lomond, Lake Maleren, Loch Ness, Lake Peijänne, Lake Peipsi, Lake Pyhäjärvi, Lake Ullswater, Lake Vänern, Lake Vesijärvi	(Kallio et al., 2015; Kutser et al., 2013)
5	China, Japan	45	Lake Biwa, Lake Erhai, Lake Suwa	(Matsushita et al., 2015)
6	New Zealand	133	Lake Brunner, Lake Hauroko, Lake Hawea, Lake Ellesmere, Lake Humuhumu, Lake Kahuparere, Lake Kai-iwi, Lake Kanono, Lake Kegonsa, Lake Mokeno, Lake Okareka, Lake Okaro, Lake Okataina, Lake Ototoa, Lake Pearson, Lake Pupuke, Lake Rotoehu, Lake Rotoiti, Lake Rotokakahi, Lake Rotokare, Lake Rotokawau, Lake Rotoma, Lake Rotonuihua, Lake Rotorua, Lake Rotorua, Lake Rototuna, Lake Sheppard, Lake Tarawera, Lake Taupo, Lake Tutuira, Lake Waikaremoana, Lake Waikere, Lake Mapourika, Lake Moeraki, Lake Monowai, Lake Paringa, Lake Te Anau, Lake Wahapo, Lake Wakatipu	(Pahlevan et al., 2020, 2022; Wang et al., 2018)
7	South Africa	26	Loskop reservoir, Theewaterskloof reservoir	(Matthews, 2014; Matthews and Bernard, 2013)
8	U.S. Mainland	487	Lake Amos, Lake BranchedOak, Lake Collins, Lake CorpusCristi, Lake Crescent, Lake DiamondPond, Lake EagleCreek, Lake EastTwin, Lake Ewell, Lake Fivemile, Flatriver reservoir, Lake Forest, Lake Fremond, Lake Fresmond, Lake Geist, Lake GingerCove, Lake GoosePond, Lake Granite, Great Salt Lake, Green Lake, Groton Lake, Hickleys Lake, Kettlebrook reservoir, Lake Attitash, Lake Champlain, Lake Hayward, Lake Manawa, Lake Mattawa, Lake Sunapee, Lake Wentworth, Lake LittleBigWood, Long Lake, Long Pond, Lowell Lake, Lake Lowerbranch, Lake Mantova, Lake Morse, Norton reservoir, Lake Okoboji, Perch Pond, Pine Lake, Pleasant Lake, Quacumquasit Pond, Quinebaug Lake, Sebago Lake, Seabasticook Lake, Silver Lake, Squam Lake, Tilden Pond, Tyler Lake, Lake Whalum, Wallum Lake	(Bradt, 2012; Dall'Omo et al., 2003, 2005; Gitelson et al., 2007; Gitelson et al., 2008; Li et al., 2013, 2015; Moore et al., 2014; Schalles, 2006; Schalles and Hladik, 2012)
9	U.S. Wisconsin	109	Big Saint Germain Lake, Big Sand Lake, Butternut Lake, Fence Lake, Fox Lake, Geneva Lake, Lac Courte Oreilles, Lac Vieux Desert, Lake Chippewa, Lake Mendota, Lake Monona, Lake Waubesa, Lake Winnebago, Metonga Lake, Pelican Lake, Rock Lake, Round Lake, Shawano Lake, Trout Lake	(Pahlevan et al., 2022, 2021, 2020)
10	U.S., Canada	198	Lake Erie, Lake Ontario, Shadow Lake, Lake Winnipeg	(Binding et al., 2008, 2010; 2011; 2013)

Appendix 2. *In situ* measurements NZ 2020

Chla water samples were filtered (using a 0.45 μm GF/F filter) and then extracted using 95% methanol (spiked with Tocopheryl acetate as an internal standard) before being sonicated and frozen to complete extraction of all chla (Hooker et al., 2009). Finally, the extract was filtered through a 0.45 μm Teflon filter. Resultant extracts were then run on an Agilent Infinity 1290 uHPLC using the method by Van Heukelem and Thomas (2001). TSM was measured gravimetrically from the dried and combusted residue on pre-combusted and pre-weighed filter pads (APHA 2540D and APHA 2540E). Laboratory triplicates of $a_{CDOM}(\lambda)$ were transferred to a 0.1 m cuvette and optical densities of the filtrates were measured for a wavelength range from 350 to 800 nm using the method described in Stedmon et al. (2000).

Radiometric measurements were made with a ship-mounted set of three hyperspectral sensors (TriOS RAMSES) installed just above the water surface. The first sensor pointed at the water surface and collected upwelling radiance $L_u(\lambda)$, which comprises both water-leaving radiance $L_w(\lambda)$ and the reflected sky irradiance $\rho_s L_s(\lambda)$. The sky irradiance $L_s(\lambda)$ was measured with a second sensor, while a third sensor measured downwelling irradiance $E_d(\lambda)$. $R_{rs}(\lambda)$ was then calculated through Simis and Olsson (2013):

$$R_{rs}(\lambda) = L_w(\lambda, 0^+)/E_d(\lambda), \quad (\text{A2})$$

$$L_w(\lambda, 0^+) = L_t(\lambda) - \rho_s L_s(\lambda). \quad (\text{A2.1})$$

Measurements were kept in this dataset if they were taken under cloud free conditions and on calm waters to ensure a correct alignment of the ship at a viewing azimuth angle (ϕ_v) $> 90^\circ$ (ideally at 135°) (Hooker and Morel, 2003). All $R_{rs}(\lambda)$ were corrected for residual reflected skylight through the method described in Jiang et al. (2020).

Appendix 3. LéXPLORE platform chla measurements

The autonomous Thetis profiler (Seabird Scientific Inc.) is deployed at the LéXPLORE platform in Lake Geneva. The profiler is comprised of hyperspectral radiometers, absorption, backscattering and attenuation sensors and a fluorescence probe for chla measurements. The top 50 m of the water column are measured by the profiler at 3-h intervals and a vertical resolution of 10 cm. A total of 278 profiles taken between October 2018 and September 2021 were used in this study. *In situ* measurements from the Thetis profiler are available in the Datalakes portal (<https://www.datalakes-eawag.ch/>).

Fluorescence probe-derived chla (CHL_F) from the Thetis profiler were measured using the WetLabs ECO Triplets BBFL2W, which has excitation and emission wavelengths at 470 and 695 nm, respectively. CHL_F estimates are impacted by non-photochemical quenching (NPQ) of fluorescence leading to an underestimation of chla in high irradiance conditions. To compensate for the underestimation, we applied the chla correction procedure developed by Roesler and Barnard (2013). We obtained a linear relationship between CHL_F and absorption line height (aLH) based chla estimates (CHL_A) during low light conditions which ensured minimal impact of NPQ to CHL_F . This linear relationship was used for daytime aLH measurements to correct NPQ-impacted CHL_F estimates. aLH was calculated using the following equations:

$$a_{BL}(\lambda_{\text{ref}}) = \frac{a(715) - a(650)}{715 - 650} * (\lambda_{\text{ref}} - 650) + a(650), a_{LH}(676) = a(676) - a_{BL}(676). \quad (\text{A3.1})$$

where a_{BL} is the absorption measured from the AC-S at the reference wavelength λ_{ref} of 676 nm. CHL_A was calculated for the upper 50 m of the water column and the top 5 m were averaged to represent *in situ* surface chla.

Appendix 4. OLCI/MSI POLYMER/C2RCC remote sensing reflectance validation

The underwater upwelling radiance (L_u) and downwelling irradiance (E_d) measurements from the Thetis profiler on LéXPLORE were measured at nadir using the Satlantic HOCR RO8W and ICSW, respectively. Both radiometers span 180 spectral channels from 300 to 1200 nm. The underwater reflectance (R_m) was calculated using the following equation:

$$R_m(\lambda, z) = \pi \left(\frac{L_u(\lambda, z)}{E_d(\lambda, z)} \right) \left[\text{dl} \right], \quad (\text{A4.1})$$

where λ is the wavelength region from 400 to 800 nm and z is depth in meters. The R_m closest to the water surface was converted to above-water remote sensing reflectance using a transmission factor following Doxaran et al. (2004):

$$R_{rs}(\lambda, z) = 0.544 * R_m(\lambda, z) \quad (\text{A4.2})$$

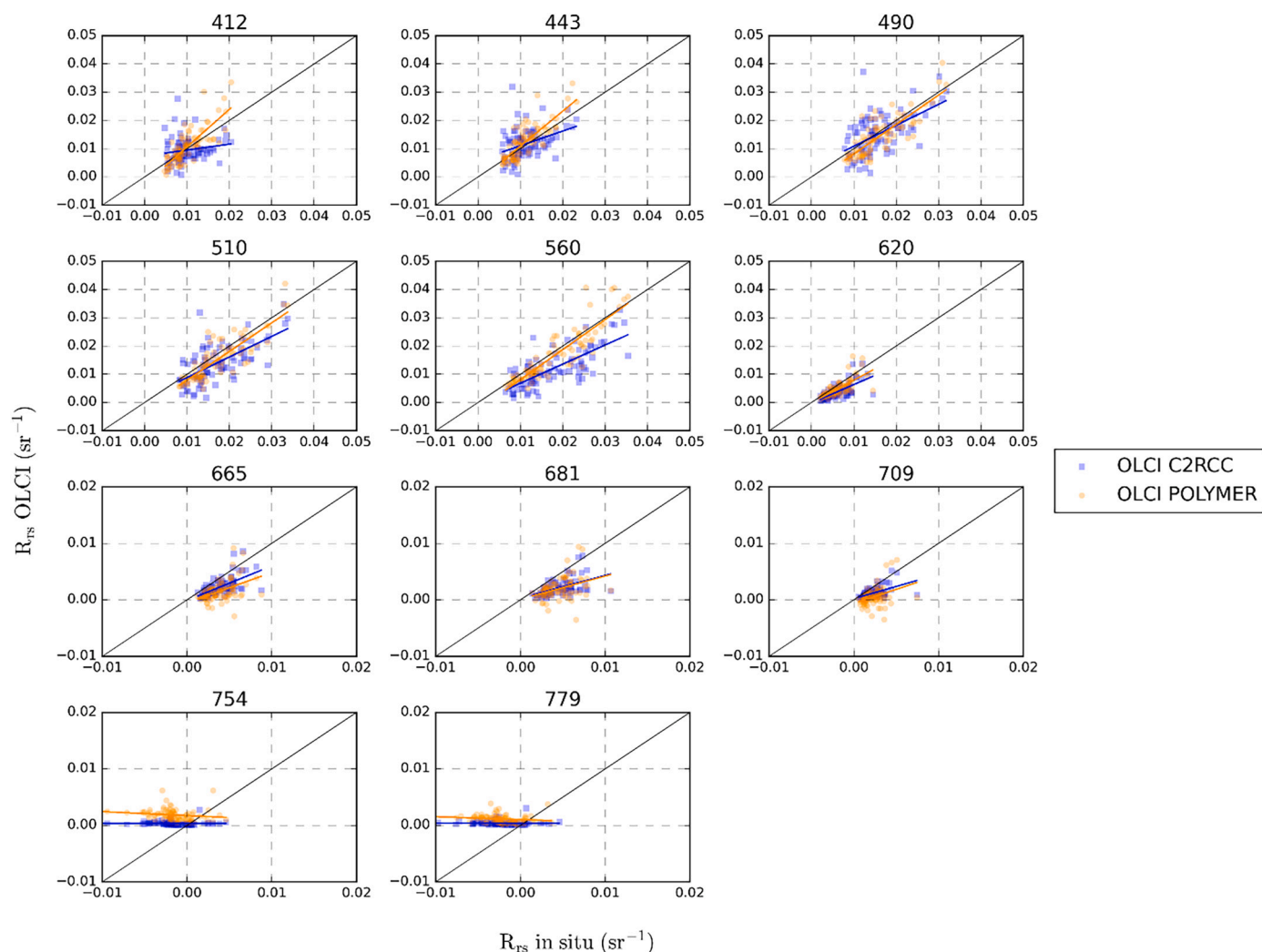


Fig. A4.1. LÉXPLORE *in situ* $R_{rs}(\lambda)$ versus OLCI C2RCC and POLYMER $R_{rs}(\lambda)$ for all available match-ups ($n = 90$). The black line represents the theoretical 1:1 regression, coloured lines linear fits through the observations. See Fig. A4.3 for the performance metrics per band. The y-axis scale changes ≥ 665 nm.

In this section, we compare the LÉXPLORE *in situ* $R_{rs}(\lambda)$ versus Sentinel-3A/B OLCI and Sentinel-2A/B MSI. The POLYMER (Steinmetz et al., 2011) and C2RCC (Brockmann et al., 2016) algorithms were used to atmospherically correct all OLCI and MSI scenes. In total, 99 OLCI match-ups with LÉXPLORE chl_a were available (see section 2.2. for a description of the match-up details). However, $R_{rs}(\lambda)$ was unavailable for nine of these OLCI chl_a match-ups due to radiometer maintenance, resulting in 90 observations available for this validation (Fig. A4.1). POLYMER estimated negative $R_{rs}(\lambda) \geq 665$ nm for 21 chl_a OLCI match-ups that were omitted from the chl_a algorithm evaluation (section 4.3) to allow for a focus on the chl_a algorithm mechanics. These negative OLCI POLYMER reflectance spectra, however, were used in this validation to investigate the AC performance. Similarly, 29 MSI chl_a match-ups were generated. 12 negative MSI POLYMER $R_{rs}(\lambda)$ were removed from the chl_a algorithm comparison but included in this $R_{rs}(\lambda)$ validation. *In situ* $R_{rs}(\lambda)$ was unavailable for three of the chl_a match-ups, yielding a total of 26 MSI match-ups for validation of AC-derived $R_{rs}(\lambda)$ (Fig. A4.2). Until 665 nm, both ACs scored the highest accuracies (Fig. A4.3). Over the blue-green bands, OLCI POLYMER was more accurate (MdSA $\leq 25\%$) than OLCI C2RCC (MdSA 10–30% higher) (Fig. A4.3).

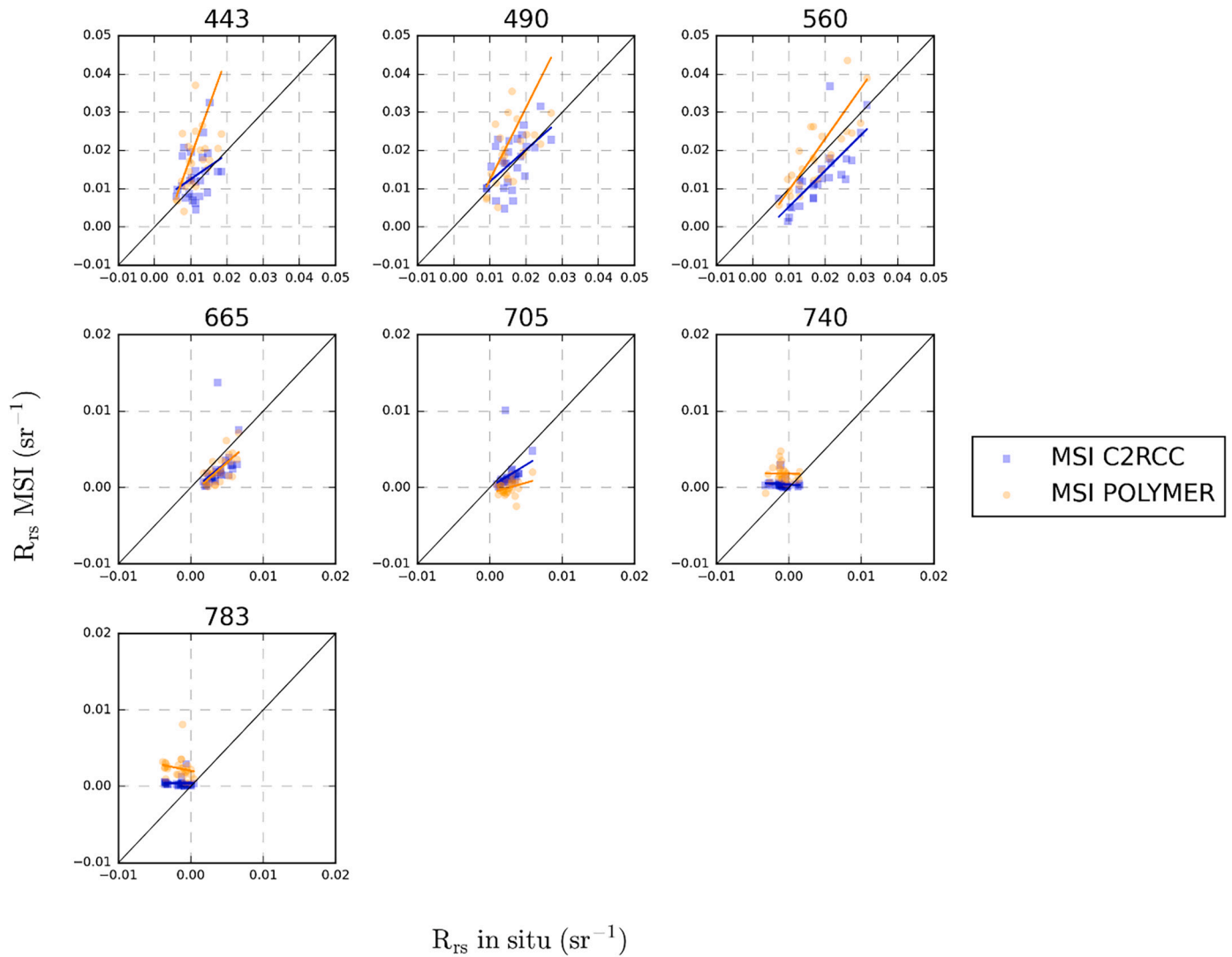


Fig. A4.2. LÉXPLORE *in situ* $R_{rs}(\lambda)$ versus MSI C2RCC and POLYMER $R_{rs}(\lambda)$ for all available match-ups ($n = 26$). See Fig. A4.1 for a description of the lines and A4.3 for performance metrics. The y-axis scale changes ≥ 665 nm.

The accuracy of the ACs in 443, 490, and 560 nm for MSI remained similar to OLCI, with minor differences in the metrics due to the smaller number of match-ups. Across the match-ups, $R_{rs}(\lambda)$ decreased sharply towards red wavelengths and reached near-zero ≥ 665 nm. Because of low signal-to-noise ratios in the red/NIR bands, *in situ* $R_{rs}(\lambda)$ became partially negative ≥ 740 nm (MSI) and 754 nm (OLCI). Similarly, POLYMER calculated a small number of negative observations ≥ 665 nm. C2RCC did not produce negative $R_{rs}(\lambda)$ because the C2RCC neural networks were originally trained using reflectance spectra with values greater than zero for wavelengths ≥ 665 nm to constrain the output (Brockmann et al., 2016). C2RCC provided more accurate $R_{rs}(\lambda)$ in these red/NIR wavelengths than POLYMER (see Fig. A4.3). The chl_a match-ups over LÉXPLORE did not include *in situ* $R_{rs}(\lambda)$ observations. Because of the negative *in situ* $R_{rs}(\lambda)$, the AC-derived $R_{rs}(\lambda)$ cannot be properly compared at 754 and 779 nm (OLCI) and 740 and 783 nm (MSI). Nonetheless, C2RCC and POLYMER values near zero appear to be feasible in these wavelengths. Because negative values obtained using POLYMER or another AC may arise for oligotrophic lakes with low signal-to-noise ratios in red/NIR, observations with negative values were set to zero in the BNNs pre-processing, preventing unpredictable behavior that would result in inflated chl_a estimates and excessive uncertainty (see section 3.2 for pre-processing details).

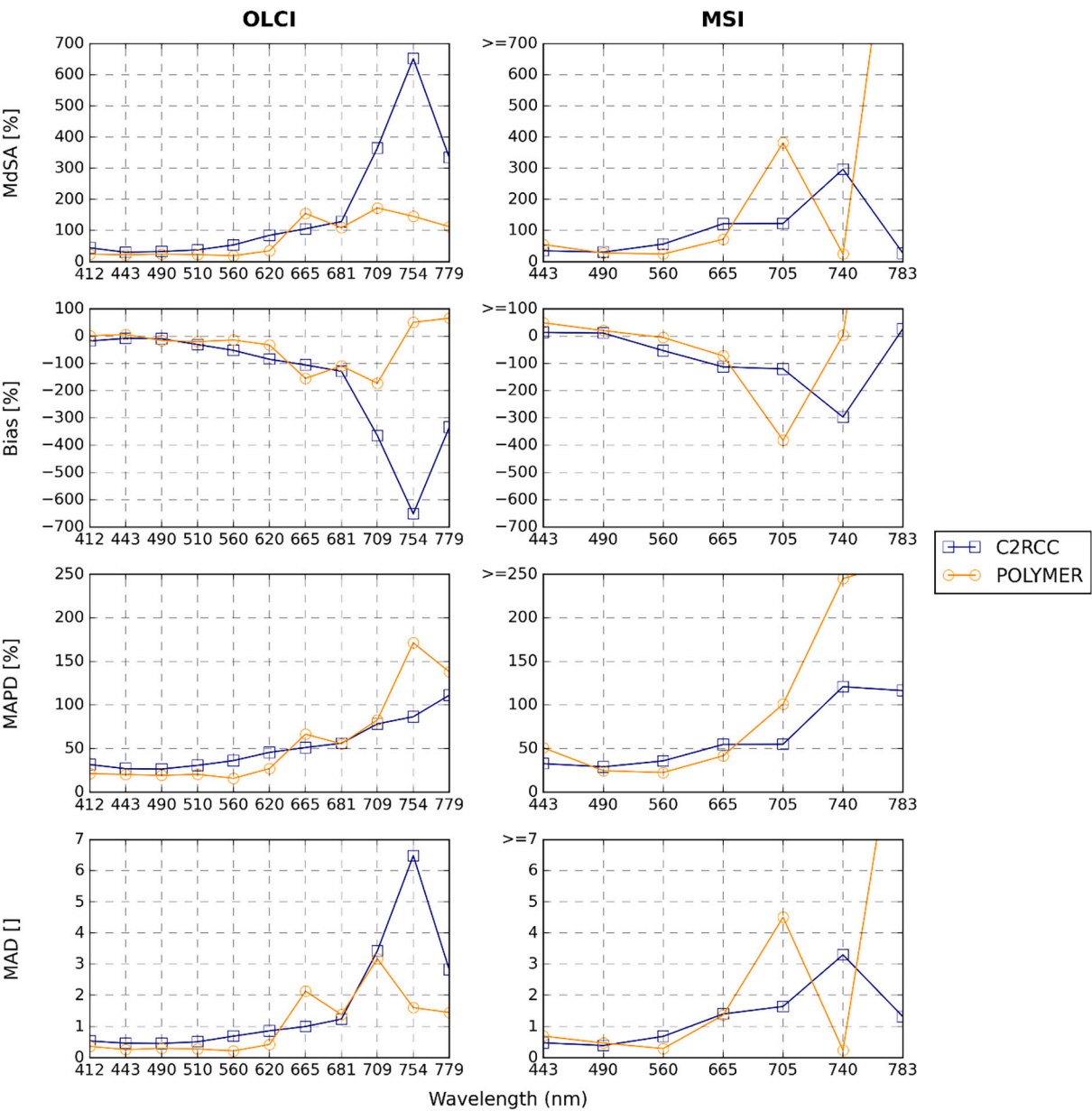


Fig. A4.3. Performance statistics per wavelength for the available OLCI and MSI match-ups generated with POLYMER and C2RCC over LÉXPLORE.

Appendix 5. BNN hyper-parameters

Hyper-parameter	Value
Training epochs	9000
Hidden layers	6
Neurons in hidden layers 1-6	200, 500, 500, 500, 200, 2
Batch-size	512
Activation function	Rectified Linear Unit (ReLU)
Dropout rate	25%
Learning rate	1.5×10^{-4}
Optimiser	Adam
Loss function	Negative-log likelihood (NLL)

Appendix 6. Optimised coefficients of reference chl_a algorithms

Algorithm	Coefficient 1	Coefficient 2	Coefficient 3	Coefficient 4	Coefficient 5
OLCI					
OC3 -opt	0.07406589	-2.02001282	0.16124464	-17.55852465	-33.62782537
OC4 -opt	0.12359851	-2.80056003	3.94756167	-23.76880782	-92.02561384
G11 -opt	-7.47220546	47.51051798	-22.15274869		
MSI					
OC3 -opt	0.07090613	-2.15824786	1.41248107	-25.01903886	-60.82627189
G11 -opt	2.22937993	38.80134518	-22.82347274		

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