

¹⁴C wiggle-matching of short tree-ring sequences from post-medieval buildings in England

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Abstract

This study tests whether accurate dating by AMS radiocarbon wiggle-matching short tree-ring series (c. 30 annual rings) in the period after AD 1510 can be achieved routinely. Such an approach has proved problematic for some intervals in the period AD 1160–1541 (Bayliss *et al.* 2017), which are before single-year calibration data are available (Stuiver 1993). We suggest that such calibration data are essential if this approach is to be employed for the informed conservation of standing buildings.

Keywords

Introduction

Over the past 25 years scientific dating has become an integral part of the processes for conservation and repair of historic buildings in England. Precise dating informs decisions about the preservation of buildings, allows us to identify significant fabric, and aids in the specification of appropriate repair strategies. Small differences in date can lead to great differences in the

significance of the extant building, and thus to great differences in the costs of the agreed solution for a particular structure.

In providing the required precise dating for historic buildings, the scientific dating method of choice is dendrochronology. In some cases, however, tree-ring analysis does not provide calendar dating, usually either because there are insufficient growth rings in the timbers or because they are of a species that is unsuitable for dendrochronology. In these cases, radiocarbon wiggle-matching is needed to provide an equivalent level of precision and reliability.

Material and Methods

Radiocarbon wiggle-matching has previously been undertaken on part of a 303-ring pine series dated by dendrochronology to AD 1367–1670 from Jermyn Street, London (Tyers *et al.* 2009). In this study, measurements from three participating AMS laboratories all provided accurate wiggle-matches from ring series covering a century or more. Insufficient data are available from this core, however, to undertake wiggle-matching on shorter sequences.

For this reason, new measurements were obtained on a 149-ring core, LED-A22, from the east principal rafter in truss 3 of one of the roofs of Ledston Hall, West Yorkshire (1.34104 W, 53.75494 N). The growth-rings in this timber span AD 1520–1668, with 20 rings of sapwood and bark edge surviving. It is part of a site sequence, LEDASQ01, which is dated by dendrochronology to AD 1424–1668 (Table 1). The ring-width data of the series are provided in Arnold *et al.* (2015).

Radiocarbon measurements were made on a total of 60 single-year tree-ring samples from this core in 2016–17. The 17 dated at the Scottish Universities Environmental Research Centre were prepared to α -cellulose, combusted, graphitised, and dated by AMS as described by Dunbar *et al.* (2016). The 18 dated at the Oxford Radiocarbon Accelerator Unit were processed using an acid-alkali-acid pretreatment followed by bleaching with sodium chlorite as described by Brock *et al.* (2010, table 1 (UW)), graphitised (Dee and Bronk Ramsey 2000), and measured by AMS (Bronk Ramsey *et al.* 2004). The 14

dated at the Bristol Radiocarbon Accelerator Mass Spectrometry Facility were processed using the base-acid-base-acid-bleach method as described by Němec *et al.* (2010), graphitised using an IonPlus AGE3 graphitisation system (Wacker *et al.* 2010), and measured using a MICADAS AMS (Synal *et al.* 2007). The 15 dated at the TUBITAK were pretreated using the acid-alkali-acid method modified from Hadjas *et al.* (2004), graphitised (Wacker *et al.* 2010) and measured by AMS on a 1 MV NEC Pelletron accelerator.

At Oxford and SUERC, $\delta^{13}\text{C}$ values, relative to VPDB, were obtained by IRMS from the gas combusted for graphitisation; at BRAMS and TUBITAK $\delta^{13}\text{C}$ values were measured by AMS.

The conventional radiocarbon ages reported for these samples, along with the rings dated, are listed in Table 2. The quoted errors are each laboratory's estimates of the total error in their dating systems. Five pairs of replicate radiocarbon measurements are available on rings dated to the same calendar year, all of which are statistically consistent at 95% confidence (Ward and Wilson 1978; Table 2). This scatter is in line with statistical expectation. Only two of the replicate $\delta^{13}\text{C}$ values are statistically consistent at 95% confidence, one is inconsistent at 95% confidence, but consistent at 99% confidence, and two are significantly different at more than 99% confidence (Ward and Wilson 1978; Table 2). These results are more scattered than would be expected on statistical grounds. This suggests that the different pre-treatment protocols used for wood samples by the Oxford and SUERC laboratories may be affecting the $\delta^{13}\text{C}$ values, especially since replicate $\delta^{13}\text{C}$ values on bone samples reported by the two laboratories in the same period are in much better agreement (Bayliss *et al.* 2016, fig. 14).

Forty-two pairs of replicate measurements, 17 groups of triplicate measurements, and one quadruple group of measurements are also available on rings dated by AMS (this study) and gas proportional counting Stuiver (1993) to the same calendar year (Figure 1). Of these 60 sets of radiocarbon ages, 52 groups are consistent at 95% confidence, 6 groups are consistent at 99% confidence, and two inconsistent at more than 99%. This scatter is rather

more than would be expected simply on the basis of statistics. Stuiver (1993) reported counting errors only, however, and, when the errors quoted for this dataset are inflated using the laboratory error multiplier suggested by Stuiver *et al.* (1998, 1045), 56 sets of measurements are consistent at 95% confidence and the remaining four at 99% confidence, which is within statistical expectation.

Wiggle-matching the entire sequence

Wiggle-matching has been undertaken using the Bayesian approach first described by Christen and Litton (1995), implemented using OxCal v4.2 (Bronk Ramsey 2009) and the IntCal13 atmospheric calibration data for the northern hemisphere (Reimer *et al.* 2013).

Figure 2 shows the model for the entire sequence of data from Ledston Hall. This has good overall agreement ($A_{comb}=76.6$, $A_n=10.5$, $n=45$; Bronk Ramsey *et al.* 2001), and estimates the final ring of the sequence to have been formed in *cal AD 1663–1671 (95% probability; SUERC-68040; Fig 2)*. This is compatible with the date of AD 1668 produced for this ring by dendrochronology (Table 1).

Wiggle-matching was then undertaken separately of the radiocarbon results quoted by each laboratory (Fig. 3a–d). The model composed of measurements made in Bristol has good overall agreement (A_{comb} : 102.8, A_n : 18.9, n : 14; Fig. 3a) and estimates that the last ring of the timber formed in *cal AD 1658–1672 (95% probability; Ring 149)*, probably in *cal AD 1662–1670 (68% probability)*. The model of measurements made at Oxford also has good overall agreement (A_{comb} : 30.9, A_n : 16.7, n : 18; Fig. 3b) and estimates the last ring of the timber to have formed in *cal AD 1655–1670 (95% probability; Ring 149)*, probably in *cal AD 1657–1664 (68% probability)*. Results from SUERC also have good overall agreement (A_{comb} : 109.9, A_n : 17.1, n : 17; Fig. 3c) and suggest that the last ring of the timber formed in *cal AD 1656–1674 (95% probability; SUERC-68040)*, probably in *cal AD 1662–1671 (68% probability)*. The wiggle match for TUBITAK measurements also

have good overall agreement (Acomb: 75.0, An: 18.3, n: 15; Fig. 3d) and suggest that the last ring of the timber formed in *cal AD 1661–1679 (95% probability; ring_149)*, probably in *cal AD 1665–1674 (68% probability)*. In all cases the estimated date of the final ring (*at 95% probability*) includes the felling date provided by dendrochronology of AD 1668.

Wiggle-matching partial sequences

Given that the length of the available tree-ring sequence is the most common limitation on successful dendrochronology in historic buildings from England, we ran a series of 25 short wiggle-matches on blocks consisting of between 29 and 30 rings. The results on the seven dated rings in each block were incorporated into a wiggle-match model that estimated the date of the final ring of the complete core. These estimates could then be compared with the known date for the final ring as derived from dendrochronology (AD 1668) to determine the accuracy of the short wiggle-matches.

The results are given in Table 3 and summarised in Figure 4. The Highest Posterior Density interval at 95% probability includes the tree-ring date for the final ring of LED-A22 (AD 1668) in all cases, and the interval at 68% probability includes the tree-ring date in all but four cases. This correspondence between the results of the wiggle-matching and the dendrochronology is greater than statistical expectation.

The long wiggle-match (AD 1160–1668)

As illustrated in Figure 2, a wiggle-match comprising the radiocarbon measurements on the 45 dated rings from Ledston Hall has good overall agreement (Acomb: 76.6; An: 10.5; n: 45) and produces posterior distributions that are compatible with the dendrochronology (AD 1668). Similarly, the wiggle-match including the results on the 18 dated rings from Jermyn Street (Tyers *et al.* 2009, fig. 4) has good overall agreement when recalculated using IntCal13 (Acomb: 35.3; An: 16.7; n: 18) and also produces posterior distributions that are compatible with the dendrochronology (AD 1670).

In contrast, the wiggle-match including the radiocarbon measurements on the 79 dated rings from the sites considered by Bayliss *et al.* (2017, fig 9) has poor overall agreement (Acomb: 1.6; An: 8.0; n: 79), and the Highest Posterior Density interval for the final ring does not include the date obtained for this ring by dendrochronology (AD 1544) even at 99% probability.

When the entire dataset is combined to form a wiggle-match sequence covering AD 1160–1668, the model has poor overall agreement (Acomb: 1.7; An: 6.4; n: 121), and the Highest Posterior Density interval for the final ring does not include the date obtained for this ring by dendrochronology (AD 1668) even at 99% probability.

Discussion

These studies in combination suggest that wiggle-matching of either short (c. 30-ring) or long (more 100-ring) tree-ring sequences produces results that are compatible with dendrochronology in the period after AD 1510. In the centuries before this, there appears to be time periods when wiggle-matching does not produce such accuracy (AD 1240–1306 and AD 1396–1532; Bayliss *et al.* 2017, table 6).

This pattern is observed when considering the datasets measured at Oxford and SUERC separately (Bayliss *et al.* 2017, table 5; Tyers *et al.* 2009, table 3; Fig. 3b–c), and so must derive from the calibration curve used, IntCal13 (Reimer *et al.* 2013).

Figures 5–6 shows the radiocarbon ages obtained on single known-age tree-rings as part of this study and those reported by Tyers *et al.* (2009) and Bayliss *et al.* (2017) in comparison to the radiocarbon ages covering this period included in IntCal13 (Reimer *et al.* 2013). The latter are on decadal samples (Wk; Hogg *et al.* 2002), single-year and decadal samples (QL; Stuiver 1993, corrected as described by Stuiver and Becker 1993; Stuiver *et al.* 1998), decadal and bi-decadal samples (UB; Hogg *et al.* 2002; Pearson *et al.* 1986), and decadal and 23-year and 24-year samples (GrN; van der Plicht *et al.* 1995).

Single-year data clearly dominate the period after AD 1510 (Fig 6), which is the period when wiggle-matching appears to be accurate within the precision quoted and the test data scatter within statistical expectation around the IntCal envelope. In the earlier period, there is more deviation between the test data and the IntCal envelope. As the statistical method of curve construction is the same in both periods (Niu *et al.* 2013), this is unlikely to be the cause of this difference, but rather there appears to be detailed structure in the atmospheric concentration of radiocarbon in the problematic periods which is not apparent from the calibration data currently available.

It should be noted that, if accurate wiggle-matching is possible in the post AD 1510 period because of the availability of single-year calibration data, then it is possible to accurately match measurements on samples of European wood against calibration data measured largely on Douglas fir trees that grew on the west coast of America. This would suggest neither intra-hemispheric locational offsets in the ^{14}C concentration of wood (McCormac *et al.* 1995) nor the translocation of ^{14}C between annual growth rings in trees (Grootes *et al.* 1989) are significant factors preventing accurate wiggle-matching.

A calibration curve based on ^{14}C measurements on single tree-rings appears to be required for wiggle-matching to provide estimates of calendar date that are accurate to within the quoted uncertainty. Such accuracy is essential if the results are to inform the long-term preservation and conservation of historic buildings.

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Table 1: Results of cross-matching site chronology LEDASQ01 with relevant independent site reference chronologies

| Reference chronology | t-value | Span of chronology | Reference |
|---|---------|--------------------|---------------------------|
| Riding House, Bolsover Castle, Derbyshire | 12.0 | AD 1494–1744 | Arnold <i>et al</i> 2005a |
| Pontefract Castle, Pontefract, West Yorkshire | 11.0 | AD 1507–1656 | Arnold <i>et al</i> 2005b |
| Little Castle, Bolsover Castle, Derbyshire | 10.4 | AD 1532–1749 | Arnold <i>et al</i> 2003 |
| Clumpcliff Farm, Rothwell, West Yorkshire | 10.2 | AD 1452–1613 | Howard <i>et al</i> 2000 |
| Auckland Castle, Bishop Auckland, County Durham | 10.1 | AD 1425–1698 | Arnold and Howard 2013 |
| Nun Appleton, Tadcaster, West Yorkshire | 9.6 | AD 1478–1657 | Arnold <i>et al</i> 2008 |

Table 2: Details of sampled tree-rings and radiocarbon results from LED-A22

| Laboratory code | Material | ¹⁴ C age (BP) | δ ¹³ C _{IRMS} (‰) | Tree-ring date (AD) |
|-----------------|--|--------------------------|---------------------------------------|---------------------|
| SUERC-68046 | <i>Quercus</i> sp. heartwood, ring 1 | 306±29 | -23.5±0.2 | 1520 |
| TUBITAK-127 | <i>Quercus</i> sp. heartwood, ring 3 | 321±29 | -24.1±1.0 | 1522 |
| OxA-34316 | <i>Quercus</i> sp. heartwood, ring 4 | 311±33 | -23.4±0.2 | 1523 |
| OxA-34317 | <i>Quercus</i> sp. heartwood, ring 4 | 357±35 | -23.8±0.2 | |
| Ring 4 | 333±25BP, T'=0.9; -23.6±0.14 ‰, T'=2.0; T'(5%)=3.8, v=1 | | | |
| BRAMS-1230 | <i>Quercus</i> sp. heartwood, ring 7 | 315±26 | -21.9±0.2 | 1526 |
| SUERC-68056 | <i>Quercus</i> sp. heartwood, ring 9 | 324±29 | -24.0±0.2 | 1528 |
| OxA-34321 | <i>Quercus</i> sp. heartwood, ring 9 | 359±33 | -24.9±0.2 | |
| Ring 9 | 339±22BP; T'=0.6; -24.5±0.14 ‰, T'=10.1; T'(5%)=3.8, v=1 | | | |
| TUBITAK-128 | <i>Quercus</i> sp. heartwood, ring 13 | 276±27 | -23.5±1.0 | 1532 |
| OxA-34275 | <i>Quercus</i> sp. heartwood, ring 14 | 262±25 | -24.9±0.2 | 1533 |
| BRAMS-1231 | <i>Quercus</i> sp. heartwood, ring 15 | 319±26 | -22.9±0.2 | 1534 |
| SUERC-68044 | <i>Quercus</i> sp. heartwood, ring 19 | 282±29 | -24.1±0.2 | 1538 |
| TUBITAK-129 | <i>Quercus</i> sp. heartwood, ring 21 | 352±50 | -24.5±1.7 | 1540 |
| BRAMS-1232 | <i>Quercus</i> sp. heartwood, ring 22 | 294±26 | -25.0± 0.2 | 1541 |
| SUERC-68054 | <i>Quercus</i> sp. heartwood, ring 24 | 315±29 | -24.2±0.2 | 1543 |
| OxA-34323 | <i>Quercus</i> sp. heartwood, ring 24 | 380±33 | -24.9±0.2 | |
| Ring 24 | 344±22BP; T'=2.2; -24.6±0.14 ‰, T'=6.1; T'(5%)=3.8, v=1 | | | |
| SUERC 68053 | <i>Quercus</i> sp. heartwood, ring 29 | 322±29 | -25.2±0.2 | 1548 |
| OxA-34281 | <i>Quercus</i> sp. heartwood, ring 29 | 255±25 | -26.2±0.2 | |
| Ring 29 | 284±19BP; T'=3.1; -25.7±0.14 ‰, T'=12.5; T'(5%)=3.8, v=1 | | | |
| BRAMS-1233 | <i>Quercus</i> sp. heartwood, ring 31 | 302±26 | -26.1±0.2 | 1550 |

| | | | | |
|-------------|---|--------|------------|------|
| OxA-34277 | <i>Quercus</i> sp. heartwood, ring 34 | 328±24 | -24.3± 0.2 | 1553 |
| TUBITAK-130 | <i>Quercus</i> sp. heartwood, ring 36 | 307±27 | -25.5±0.8 | 1555 |
| SUERC-68051 | <i>Quercus</i> sp. heartwood, ring 39 | 335±29 | -24.6±0.2 | 1558 |
| OxA-34319 | <i>Quercus</i> sp. heartwood, ring 44 | 334±32 | -25.4±0.2 | 1563 |
| TUBITAK-131 | <i>Quercus</i> sp. heartwood, ring 45 | 310±27 | -25.3±0.8 | 1564 |
| BRAMS-1234 | <i>Quercus</i> sp. heartwood, ring 47 | 353±26 | -22.4±0.2 | 1566 |
| SUERC-68034 | <i>Quercus</i> sp. heartwood, ring 49 | 306±29 | -25.1±0.2 | 1568 |
| OxA-34457 | <i>Quercus</i> sp. heartwood, ring 54 | 364±27 | -25.9±0.2 | 1573 |
| BRAMS-1235 | <i>Quercus</i> sp. heartwood, ring 56 | 382±26 | -28.4±0.2 | 1575 |
| TUBITAK-132 | <i>Quercus</i> sp. heartwood, ring 58 | 290±48 | -25.1±1.5 | 1577 |
| SUERC-68035 | <i>Quercus</i> sp. heartwood, ring 59 | 346±29 | -25.8±0.2 | 1578 |
| TUBITAK-133 | <i>Quercus</i> sp. heartwood, ring 62 | 286±27 | -26.7±0.8 | 1581 |
| OxA-34279 | <i>Quercus</i> sp. heartwood, ring 64 | 331±26 | -25.9±0.2 | 1583 |
| BRAMS-1236 | <i>Quercus</i> sp. heartwood, ring 67 | 336±26 | -25.8±0.2 | 1586 |
| SUERC-68052 | <i>Quercus</i> sp. heartwood, ring 69 | 310±29 | -25.9±0.2 | 1588 |
| OxA-34282 | <i>Quercus</i> sp. heartwood, ring 74 | 297±26 | -27.1±0.2 | 1593 |
| TUBITAK-134 | <i>Quercus</i> sp. heartwood, ring 77 | 396±29 | -26.3±0.9 | 1596 |
| SUERC-68036 | <i>Quercus</i> sp. heartwood, ring 79 | 334±29 | -25.6±0.2 | 1598 |
| OxA-34283 | <i>Quercus</i> sp. heartwood, ring 84 | 336±26 | -26.4±0.2 | 1603 |
| TUBITAK-135 | <i>Quercus</i> sp. heartwood, ring 85 | 336±28 | -24.9±0.9 | 1604 |
| BRAMS-1238 | <i>Quercus</i> sp. heartwood, ring 86 | 365±26 | -24.4± 0.2 | 1605 |
| SUERC-68050 | <i>Quercus</i> sp. heartwood, ring 89 | 347±29 | -25.2± 0.2 | 1608 |
| OxA-34278 | <i>Quercus</i> sp. heartwood, ring 94 | 341±25 | -24.8±0.2 | 1613 |
| TUBITAK-136 | <i>Quercus</i> sp. heartwood, ring 96 | 416±42 | -26.3±0.9 | 1615 |
| BRAMS-1239 | <i>Quercus</i> sp. heartwood, ring 98 | 370±26 | -23.1±0.2 | 1617 |
| SUERC-68042 | <i>Quercus</i> sp. heartwood, ring 99 | 364±29 | -23.7±0.2 | 1618 |
| OxA-34318 | <i>Quercus</i> sp. heartwood, ring 104 | 374±31 | -24.1± 0.2 | 1623 |
| TUBITAK-137 | <i>Quercus</i> sp. heartwood, ring 107 | 351±27 | -23.0±0.7 | 1626 |
| SUERC-68055 | <i>Quercus</i> sp. heartwood, ring 109 | 313±29 | -23.8±0.2 | 1628 |
| OxA-34322 | <i>Quercus</i> sp. heartwood, ring 109 | 351±34 | -24.1±0.2 | |
| Ring 109 | 284±19BP; T'=3.1; -24.0±0.14 ‰, T'=1.1; T'(5%)=3.8, v=1 | | | |
| BRAMS-1240 | <i>Quercus</i> sp. heartwood, ring 110 | 348±26 | -23.3±0.2 | 1629 |
| OxA-34280 | <i>Quercus</i> sp. heartwood, ring 114 | 321±26 | -25.0±0.2 | 1633 |
| TUBITAK-138 | <i>Quercus</i> sp. heartwood, ring 115 | 312±26 | -22.7±0.7 | 1634 |
| SUERC-68043 | <i>Quercus</i> sp. heartwood, ring 119 | 320±29 | -22.9±0.2 | 1638 |
| BRAMS-1241 | <i>Quercus</i> sp. heartwood, ring 120 | 291±26 | -22.7±0.2 | 1639 |
| TUBITAK-139 | <i>Quercus</i> sp. heartwood, ring 122 | 271±28 | -24.0±0.9 | 1641 |
| OxA-34284 | <i>Quercus</i> sp. heartwood, ring 124 | 275±25 | -25.7±0.2 | 1643 |
| BRAMS-1242 | <i>Quercus</i> sp. heartwood, ring 125 | 263±26 | -25.7±0.2 | 1644 |
| SUERC-68041 | <i>Quercus</i> sp. heartwood, ring 129 | 214±29 | -23.7±0.2 | 1648 |

| | | | | |
|-------------|--|--------|-----------|------|
| BRAMS-1243 | <i>Quercus</i> sp. heartwood, ring 131 | 301±26 | -21.7±0.2 | 1650 |
| OxA-34320 | <i>Quercus</i> sp. heartwood, ring 134 | 292±32 | -24.3±0.2 | 1653 |
| TUBITAK-140 | <i>Quercus</i> sp. heartwood, ring 135 | 212±29 | -23.8±1.1 | 1654 |
| SUERC-68045 | <i>Quercus</i> sp. heartwood, ring 139 | 267±29 | -23.4±0.2 | 1658 |
| OxA-34276 | <i>Quercus</i> sp. heartwood, ring 144 | 214±24 | -24.2±0.2 | 1663 |
| TUBITAK-141 | <i>Quercus</i> sp. heartwood, ring 146 | 248±55 | -26.3±2.3 | 1665 |
| BRAMS-1244 | <i>Quercus</i> sp. heartwood, ring 148 | 214±26 | -21.9±0.2 | 1667 |
| SUERC-68040 | <i>Quercus</i> sp. heartwood, ring 149 | 222±29 | -24.0±0.2 | 1668 |

Table 3: Summary of the estimated dates for the final ring of LED-A22 (dated by dendrochronology to AD 1668) from wiggle-matching 29–30-year blocks (An: 26.7, n: 7 for all).

| Rings | Acomb | <i>Highest Posterior Density interval (cal AD)</i> | |
|---------|-------|--|---|
| | | <i>68% probability</i> | <i>95% probability</i> |
| 1–29 | 48.0 | 1660–1676 (59%) or 1683–1688 (9%) | 1655–1709 (89%) or 1760–1770 (6%) |
| 4–34 | 51.5 | 1659–1675 | 1655–1693 (93%) or 1700–1706 (2%) |
| 9–39 | 50.8 | 1660–1675 (60%) or 1683–1687 (8%) | 1655–1706 |
| 14–44 | 63.2 | 1664–1676 | 1657–1694 |
| 19–49 | 83.3 | 1655–1689 | 1645–1699 |
| 24–54 | 72.3 | 1650–1657 (10%) or 1665–1694 (58%) | 1643–1706 |
| 29–59 | 90.6 | 1653–1678 (58%) or 1686–1692 (10%) | 1647–1700 |
| 34–64 | 102.7 | 1655–1704 | 1596–1620 (11%) or 1634–1723 (84%) |
| 39–69 | 95.1 | 1599–1615 (20%) or 1651–1682 (38%) or 1697–1704 (8%) or 1719–1721 (2%) | 1593–1636 (27%) or 1642–1725 (68%) |
| 44–74 | 95.7 | 1598–1617 (42%) or 1662–1670 (9%) or 1709–1722 (17%) | 1595–1631 (44%) or 1640–1677 (26%) or 1695–1724 (25%) |
| 49–79 | 91.5 | 1599–1617 (39%) or 1649–1670 (29%) | 1592–1675 (80%) or 1694–1715 (15%) |
| 54–84 | 98.4 | 1597–1618 (43%) or 1651–1668 (16%) or 1698–1707 (9%) | 1576–1675 (81%) or 1692–1711 (14%) |
| 59–89 | 103.6 | 1605–1613 (8%) or 1627–1668 (60%) | 1576–1586 (2%) or 1596–1677 (92%) or 1696–1702 (1%) |
| 64–94 | 118.4 | 1621–1659 | 1601–1676 |
| 69–99 | 116.8 | 1623–1659 | 1607–1672 |
| 74–104 | 114.1 | 1637–1664 | 1563–1569 (1%) or 1611–1672 (94%) |
| 79–109 | 117.0 | 1555–1567 (18%) or 1637–1667 (50%) | 1544–1573 (25%) or 1617–1675 (70%) |
| 84–114 | 115.5 | 1551–1568 (31%) or 1649–1668 (37%) | 1538–1573 (37%) or 1616–1677 (58%) |
| 89–119 | 117.5 | 1548–1568 (35%) or 1650–1670 (33%) | 1529–1574 (46%) or 1609–1635 (11%) or 1646–1675 (38%) |
| 94–124 | 124.8 | 1553–1564 (36%) or 1658–1670 (32%) | 1546–1574 (50%) or 1649–1677 (45%) |
| 99–129 | 92.7 | 1664–1674 | 1555–1557 (1%) or 1657–1679 (94%) |
| 104–134 | 78.3 | 1662–1674 | 1552–1563 (6%) or 1656–1679 (89%) |
| 109–139 | 97.7 | 1664–1673 | 1656–1678 |
| 114–144 | 104.1 | 1665–1673 | 1659–1675 |
| 119–149 | 98.8 | 1664–1671 | 1659–1675 |

Figure 1: Offsets between radiocarbon ages on single tree-ring from this study and measurements on single tree-rings of the same calendar date reported by Stuiver (1993, corrected as described by Stuiver and Becker 1993).

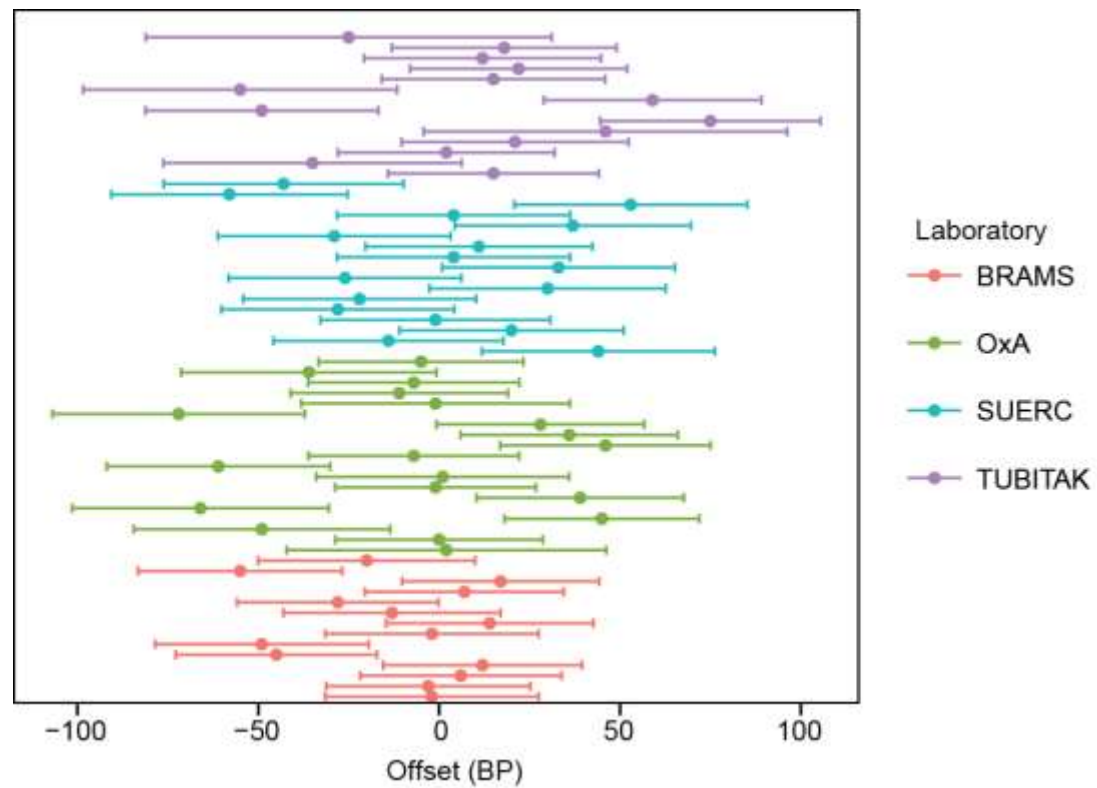
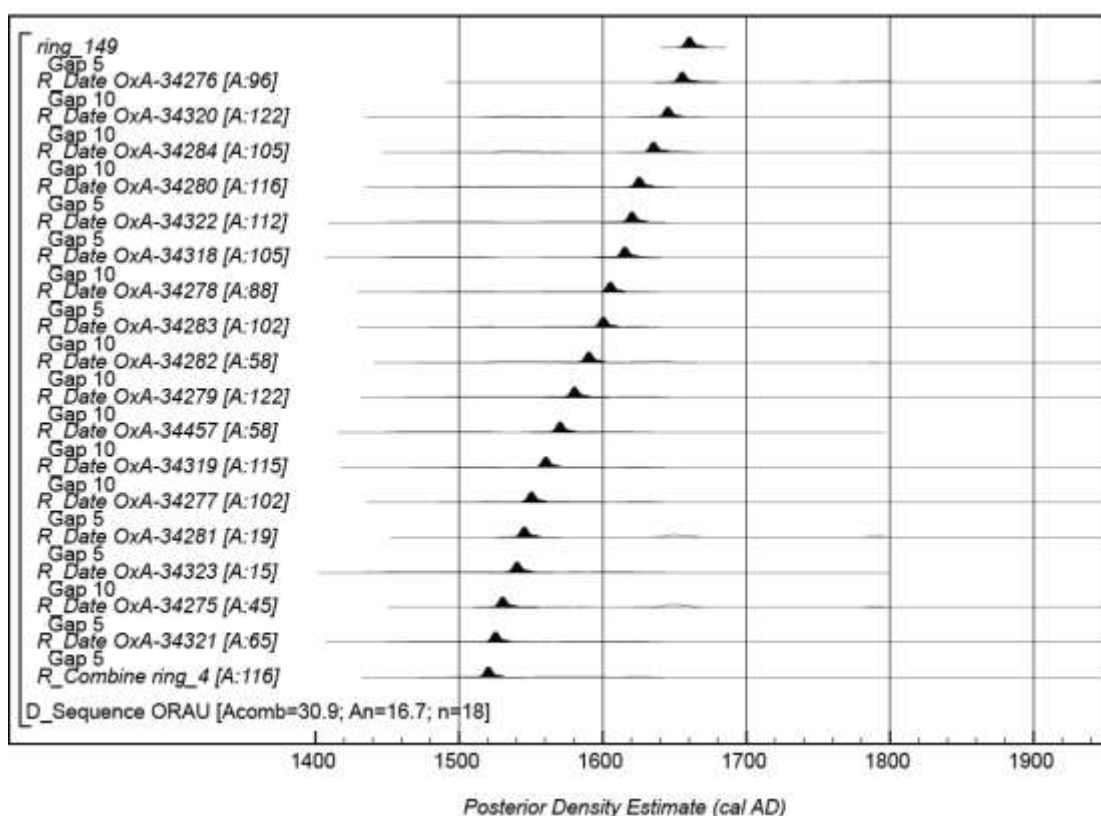
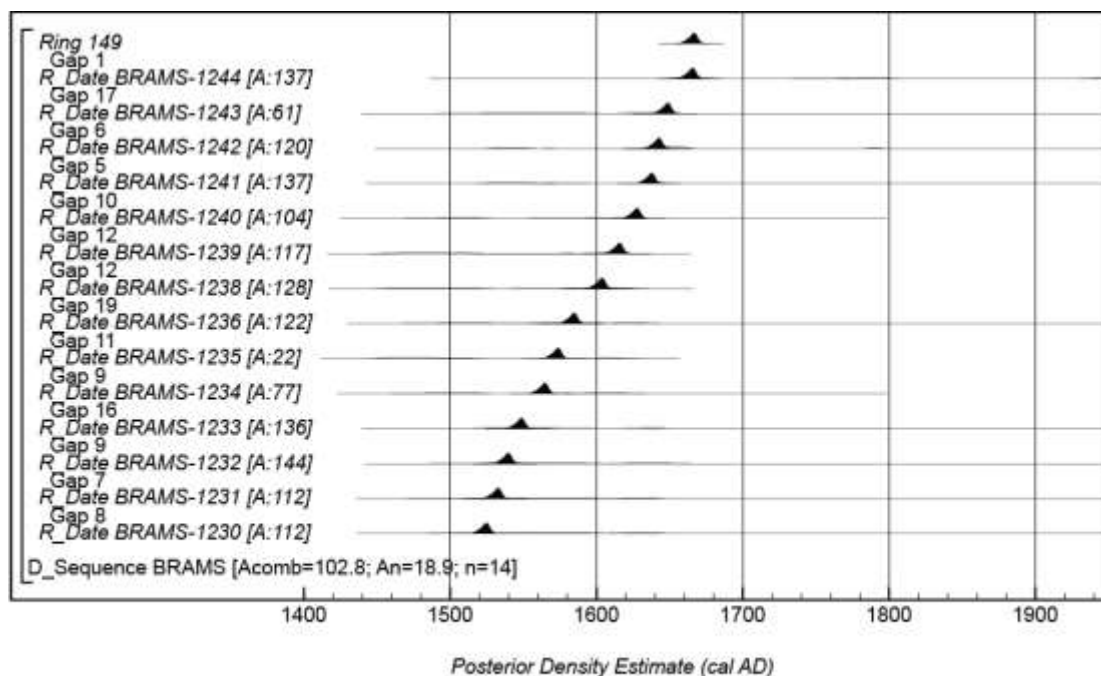


Figure 2: Probability distributions of dates from LED-A22. Each distribution represents the relative probability that an event occurs at a particular time. For each of the dates two distributions have been plotted: one in outline, which is the result of simple radiocarbon calibration, and a solid one, based on the wiggle-match sequence. Distributions other than those relating to particular samples, correspond to aspects of the model. For example, the distribution 'SUERC-68040' is the estimated date of the final ring of this core. The large square brackets down the left-hand side of the diagram along with the CQL2 keywords (Bronk Ramsey 2009) define the model exactly.



Figure 3: Probability distributions of dates from LED-A22 (a) BRAMS-, (b) OxA-, (c) SUERC, and (d) TUBITAK. The format is identical to that of Figure 2. The large square brackets down the left-hand side of the diagram along with the CQL2 keywords define the model exactly



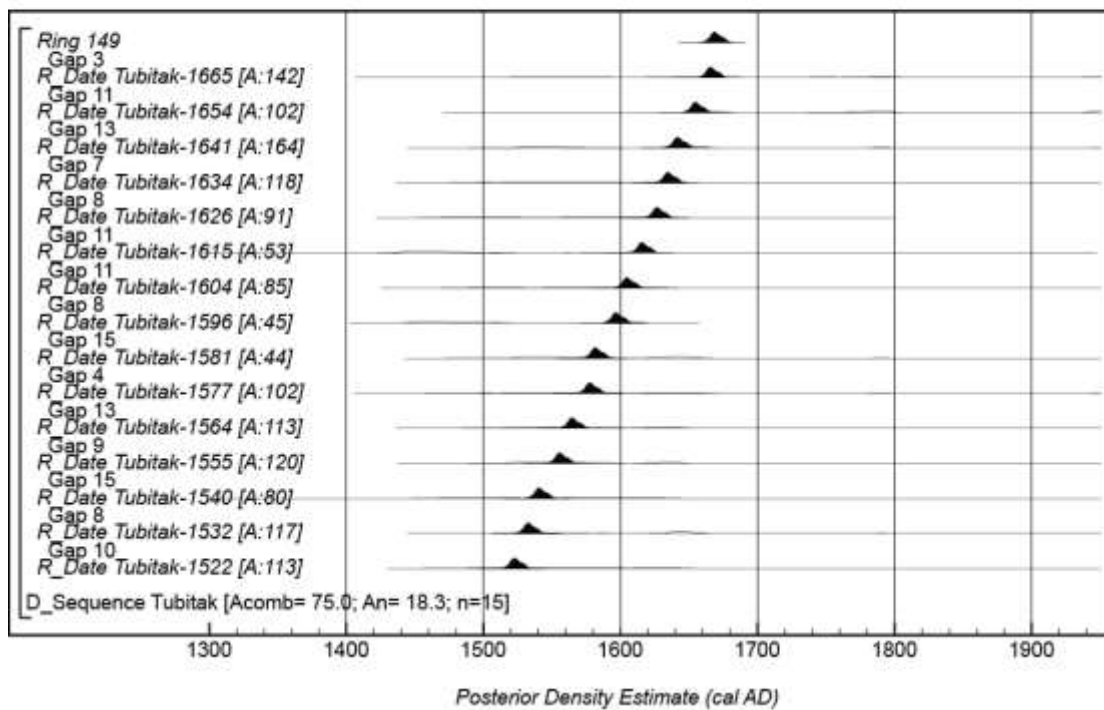
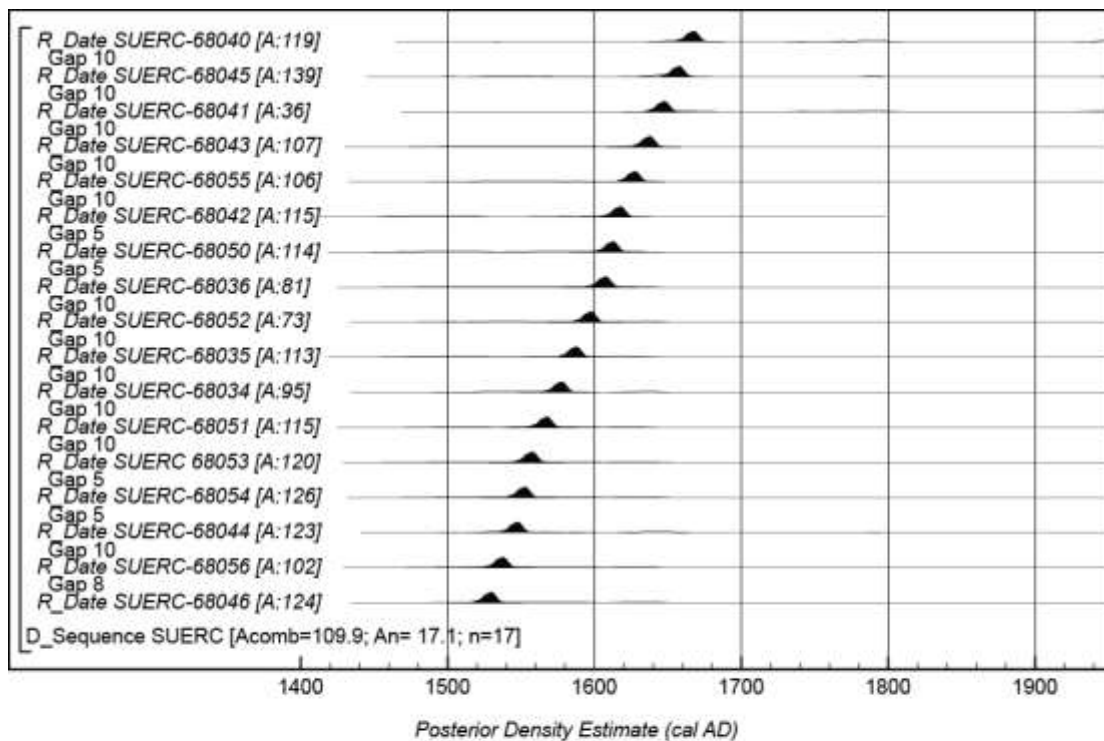


Figure 4: Posterior density estimates for the final ring of LED-A22, derived from the short wiggle-matches based on sequences of 29–30 tree-rings.

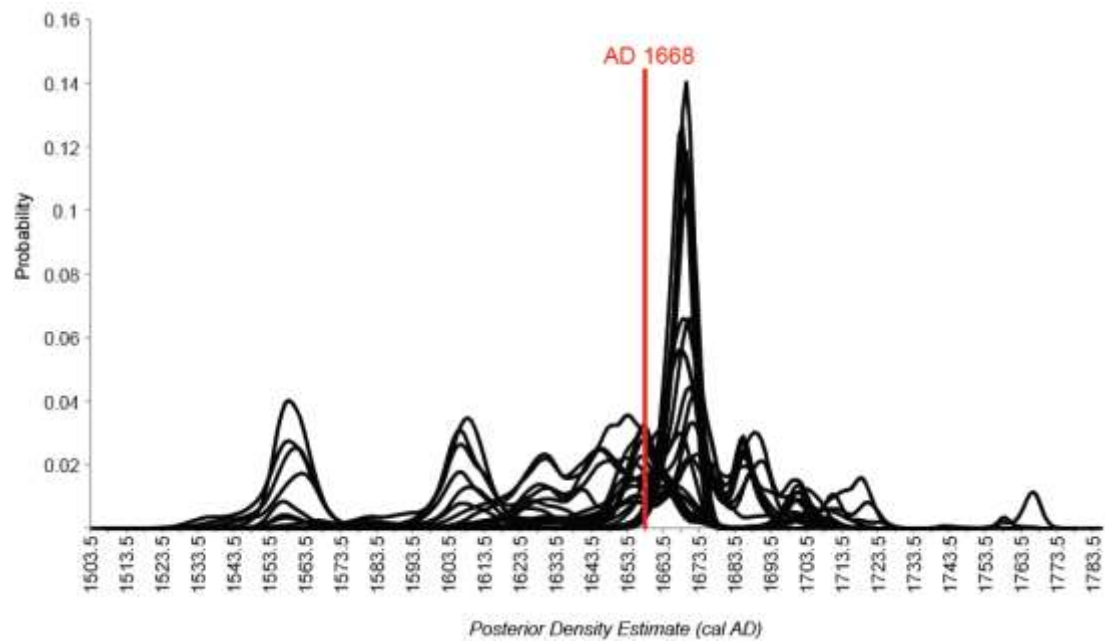


Figure 5: Radiocarbon ages of known-age tree-ring rings AD 1150–1668: single years (OxA-, SUERC-, GrA-, BRAMS-, TUBITAK-; this study, Bayliss *et al.* 2017, Tyers *et al.* 2009), decadal samples (Wk; Hogg *et al.* 2002), single-year and decadal samples (QL; Stuiver 1993 as corrected by Stuiver and Becker 1993; Stuiver *et al.* 1998), decadal and bi-decadal samples (UB; Hogg *et al.* 2002; Pearson *et al.* 1986), decadal and 23-year and 24-year samples (GrN: van der Plicht *et al.* 1995). Inset shows detail of period AD 1510–1670.

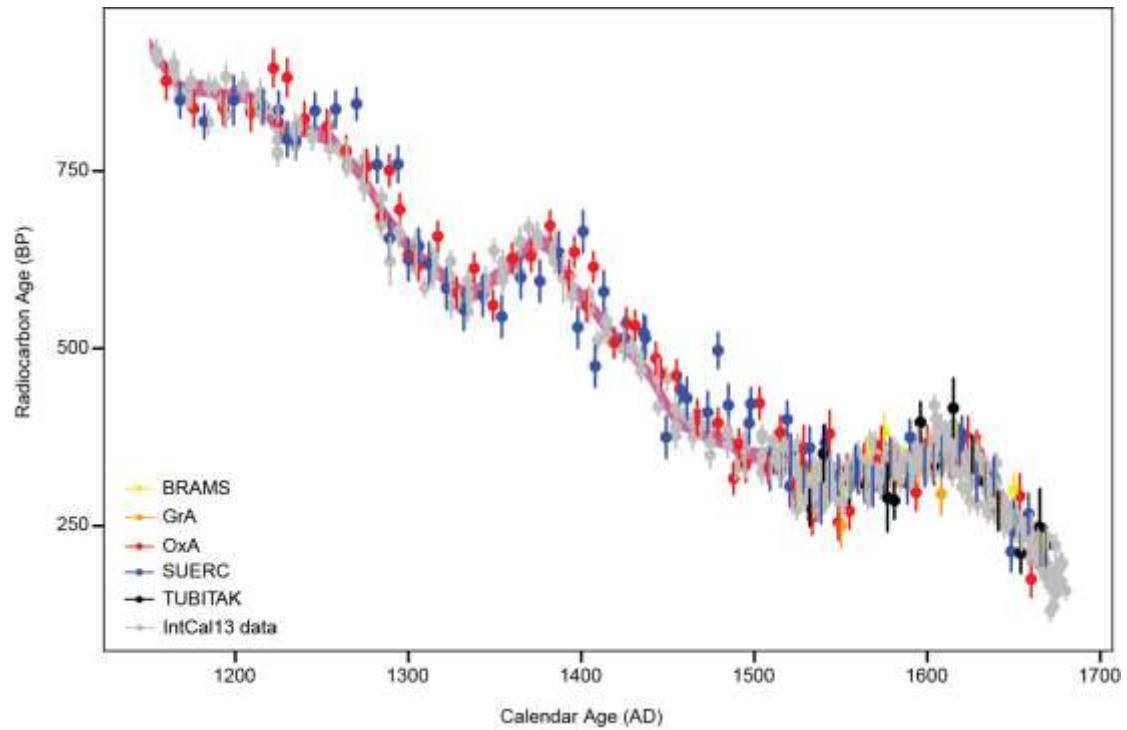


Figure 6: Radiocarbon ages of known-age tree-ring rings AD 1510–1670: single years (OxA-, SUERC-, GrA-, BRAMS-, TUBITAK-; this study, Bayliss *et al.* 2017, Tyers *et al.* 2009), decadal samples (Wk; Hogg *et al.* 2002), single-year samples (QL; Stuiver 1993 as corrected by Stuiver and Becker 1993; Stuiver *et al.* 1998), decadal and bi-decadal samples (UB; Hogg *et al.* 2002; Pearson *et al.* 1986), decadal and 23-year and 24-year samples (GrN: van der Plicht *et al.* 1995).

